Isotopic Composition and sources of Organic Carbon Pools within the Tana River Basin, (Kenya).

Fredrick Tamooh ^{1,2}, Karel Van Den Meersche ^{3,4}, Alberto Borges ⁵, Roel Merckx ¹, Frank Dehairs ³, Filip Meysman ^{3,4} & Steven Bouillon ¹

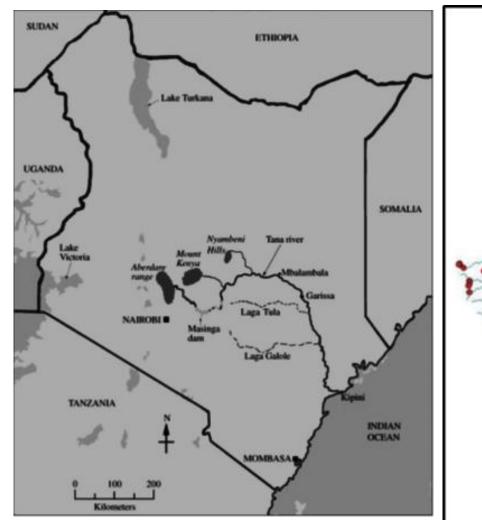
(1) Katholieke Universiteit Leuven, dept. of Earth & Environmental Sciences, Belgium; E-mail : fredrick.tamooh@ees.kuleuven.be
(2) Kenya Wildlife Services, P.O. Box 82144-80100, Mombasa, Kenya (3) Department of Analytical and Environmental Chemistry, Vrije Universiteit Brussel, Belgium (4) Netherlands Institute of Ecology (NIOO-CEME), Yerseke, The Netherlands (5) Unité d'Océanographie Chimique Université de Liège, Belgium

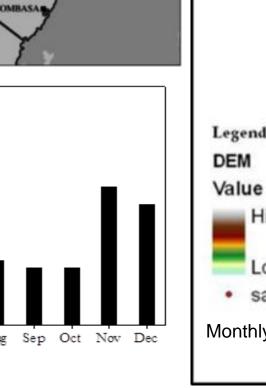
Introduction

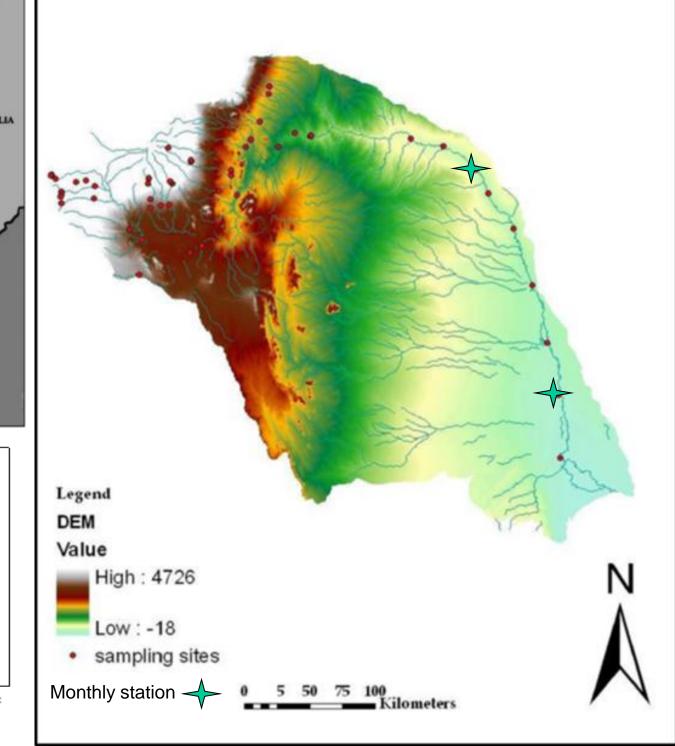
- ➤ Rivers play an important role in the global carbon cycle, and process ~1.9 Pg C annually (Cole et al., 2007). Rivers do not merely transport carbon from the terrestrial to the oceanic environment, but also bury and process organic matter, typically acting as a source of CO₂ to the atmosphere (Cole and Caraco 2001, Mayorga et al,. 2005).
- It is critical to understand carbon cycling both on a global and a watershed scale. However, there are few studies which quantify carbon fluxes in tropical rivers, and data for the African continent are particularly scarce.
- In this study, we report the altitudinal and seasonal patterns in carbon pools and their stable isotope compositions in Tana River Basin (Kenya)

Site and Methods

- ➤ The Tana River is the longest river system in Kenya (~1300 km), with a catchment area of ~130,000 km² (Kitheka et al., 2005).
- ➤ The main perennial source areas of the river are Mount Kenya (up to 5199 m asl), the Abardares ranges in the central highlands of Kenya, and the Nyambene Hills in eastern Kenya.
- The basin in general experiences a bimodal rainfall pattern: long rains between March and May and short rains between October and December.
- > Data from field campaigns throughout the river basin are presented from three campaigns in February 2008 (Bouillon et al., 2009; dry-season), September to November 2009 (wet-season) and June-July 2010 (end-of-wet-season).
- Furthermore, monthly sampling was initiated in January 2009 at several locations (ongoing), and data up to March 2010 are presented here for 2 of the downstream sites (Garissa and Tana River Primate Reserve). Extensive flood plains are located between these 2 locations, flooding is irregular due to regulation of river flows by reservoirs upstream.
- The samples for total suspended matter (TSM) were filtered through pre-combusted and pre-weighed, 47-mm-diameter Whatman GF/F filters, dried and re-weighed, while samples for POC and δ¹³C-POC were filtered on pre-combusted 25 mm Whatman GF/F filters, acidified, dried and packed in Ag cups. Soil and sediments samples were collected from all sampling sites, subsamples grounded, decarbonated and similarly packed in Ag cups. POC, δ¹³C-POC, soil and sediments were measured with standard techniques (EA-IRMS). DOC and δ¹³C-DOC samples were measured with a TOC analyzer coupled to a Thermo DeltaPlus IRMS.







Conclusions

phytoplankton production.

> δ¹³C constrained from organic matter

➤ High POC:Chl a ratios suggest

> TSM and POC delivery is episodic

> In the lower section of the Tana River,

mobilized at intermediate altitudes

during peak discharge, and mostly

POC mainly originates from areas with

a significant contribution by C4 plant

species during high discharge, while

during low discharge POC is

predominantly derived from C3 plant

dominate the riverine DOC and POC

negligible contribution from in-stream

C3 derived organic matter

Figure 1:Location of the Tana River basin, sampling locations and monthly discharge.

pools.

Results & Discussion

Total suspended matter and particulate organic Carbon

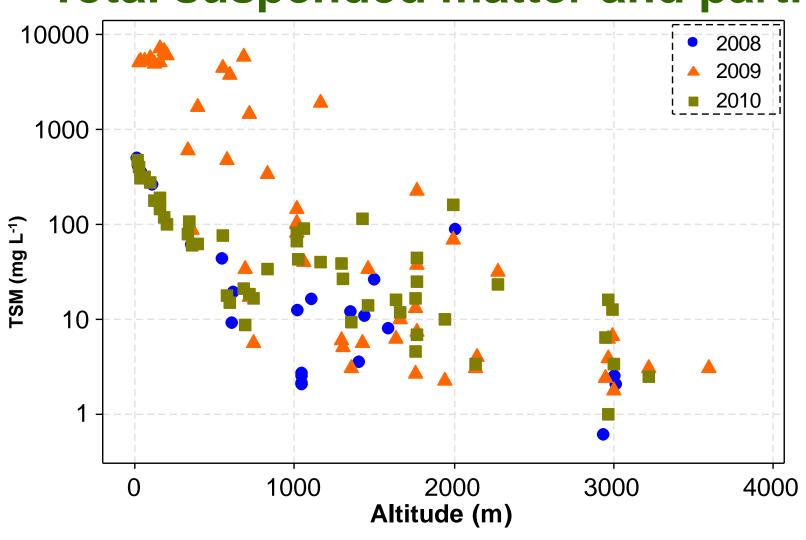


Figure 2: Altitudinal profile of TSM

- ➤ A consistent downstream increase in TSM was observed during all three sampling campaigns.
- TSM values were similar for the dry-season and end-of-wet-season datasets (p>0.05), but significantly higher during the wet-season campaign.

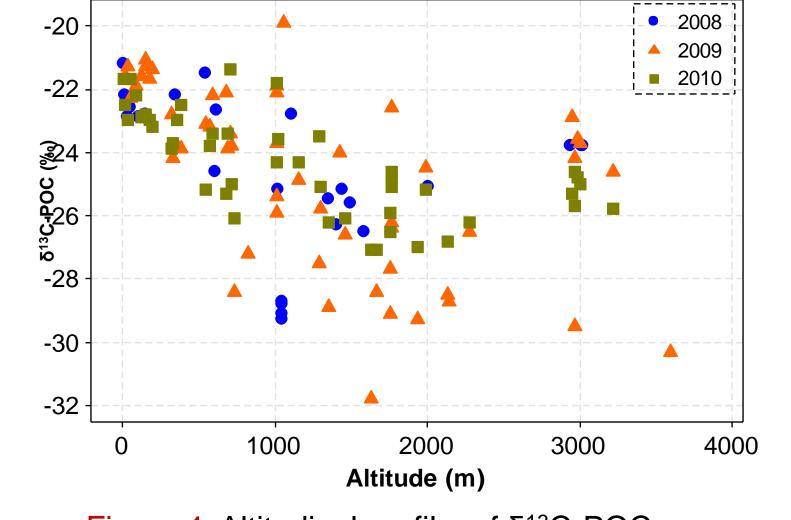


Figure 4: Altitudinal profile of δ¹³C-POC

20 Sediment

15
0 500 1000 1500 2000 2500 3000 3500

Altitude (m)

Figure 6: Altitudinal profile of soil and sediment % OC

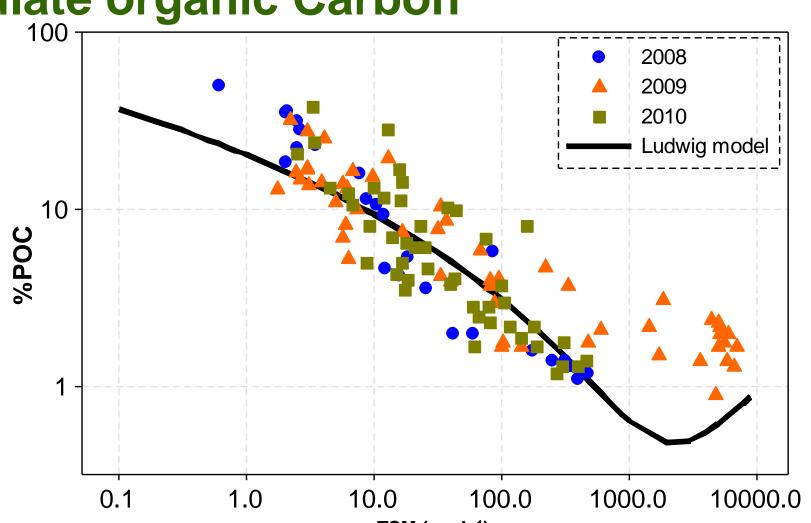


Figure 3: Comparison of % POC and TSM along Tana River Basin and Ludwig et al., 1996 Model

- The POC concentrations shows similar trends as those in TSM, i.e. a consistent downstream increase during all sampling campaigns (p<0.01).
- TSM & %POC followed the classical inverse relationship for all seasons sampled i.e. dilution of %POC with increase in TSM (p<0.01).

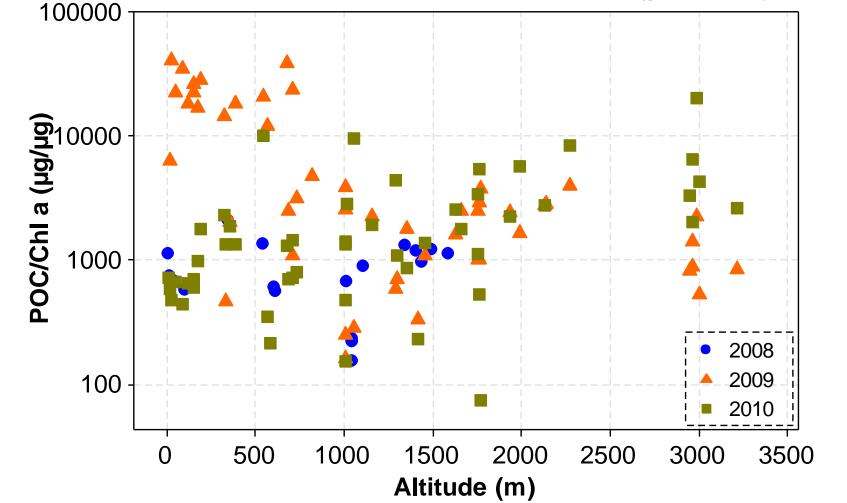
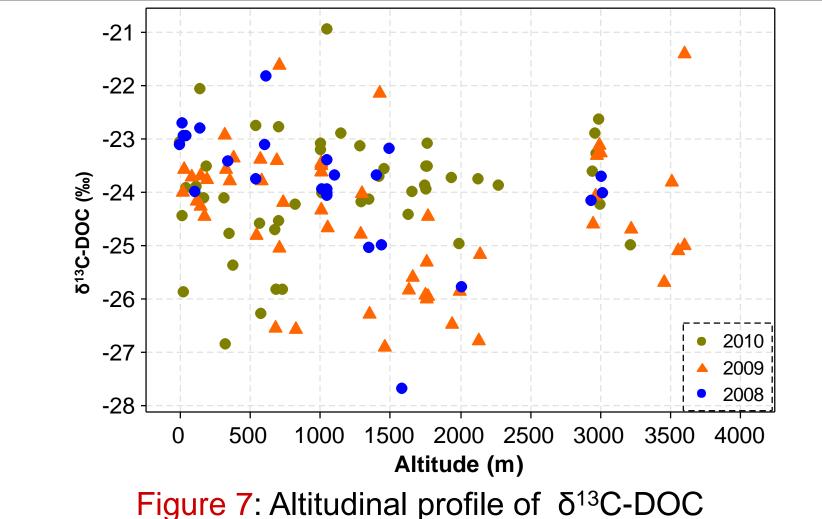


Figure 5: Altitudinal profile of POC/Chl a ratios

- ➤ The %POC/TSM relationship fits well with empirical Model based on Ludwig et al. (1996) data.
- > δ¹³C-POC increased downstream during all three sampling campaigns (p<0.01), and were predominantly of terrestrial origin as reflected by generally high POC/ChI a ratios. This trend thus reflects an increasing contribution of C4-derived carbon downstream. However, different seasons were not significantly different (p>0.05).
- ➤ Soil and sediment %OC decreased consistently downstream (p<0.01) due to minimal organic matter decomposition in high altitude for soil. Sediment had less %OC compared to soil due to large particle size.



6 2008 2009 2010 3 2010 DOC dominates POC dominates 0

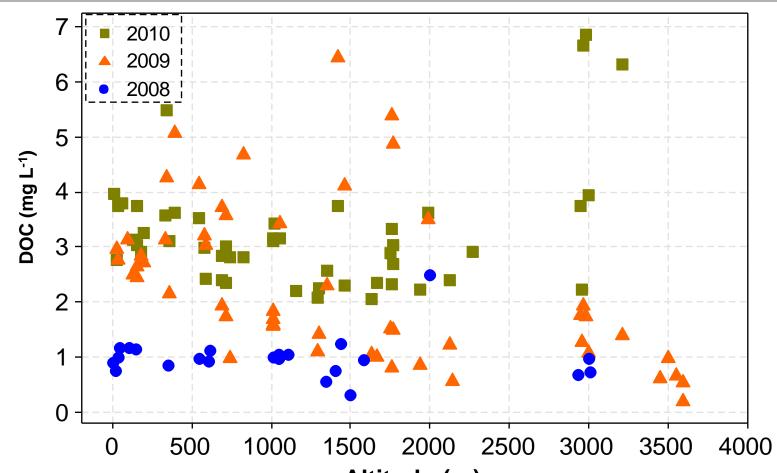
Figure 9: Relation between DOC:POC ratios & TSM.

Seasonal variation in organic composition

Garissa

- ◆ - TRPR

Discharge



Altitude (m)
Figure 8: Altitudinal Profile of DOC

- > DOC concentration was higher in wetseason due to flushing effect of organic.
- The range of δ^{13} C-DOC values (-27.7 to -20.9‰) suggests the source of DOC is predominantly of C3 origin and minimal or no in-stream autochthonous production .
- The DOC:POC ratios show a significant inverse relationship with TSM typical of an erosive riverine system.

References

species.

Bouillon S , Abril G, Borges AV, Dehairs F, Govers G, Hughes H.J, Merckx R, Meysman F, Nyunja J, Osburn C, & Middelburg J.J. *Biogeosciences* 6: 2475-2493

Cole, J. J., Prairie, Y. T., Caraco, 5 N. F., McDowell, W. H., Tranvik, L. J., Striegl, R. G., Duarte, C. M., Kortelainen, P., Downing, J. A., Middelburg, J. J., and Melack, J *Ecosystems*,10, 171–184, (2007).

Cole, J. J. and Caraco, N. F., *Mar. Freshwater Res.* 52, 101–110 (2001),

Kitheka, J.U, Obiero, M., and Nthenge, P. *Estuar. Coast. Shelf Sci.*, 63, 455–468 (2005).

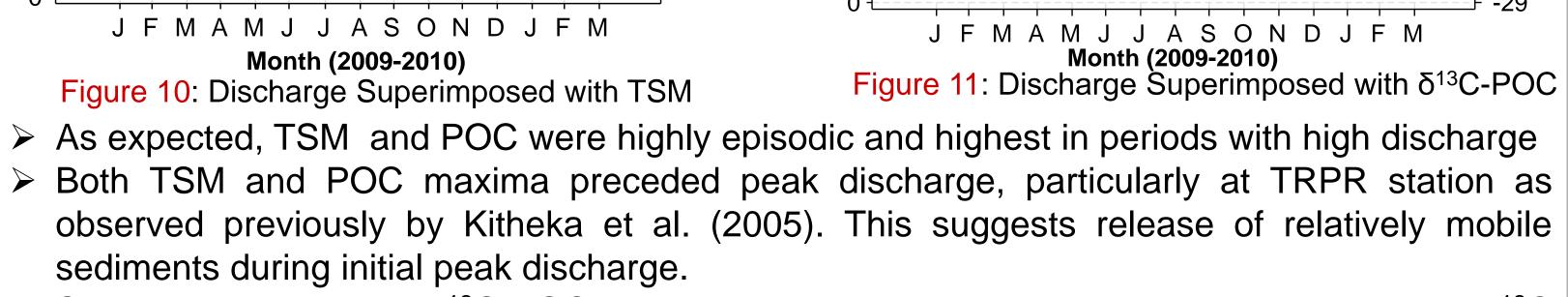
Ludwig, W., Probst, J.L., Kempe, S. Global Biogeochemical Cycles 10, 23-41 (1996).

Mayorga, E., Aufdenkampe, A. K., Masiello, C. A., Krusche, A. V., Hedges, J. I., Quay, P. D., Richey, J. E., and Brown, T. A. Nature, 436, 538–541 (2005).

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= 200

3000 g

2000**°**

Seasonal patterns in δ^{13} C-POC signatures at both stations coincided closely, with δ^{13} C increasing markedly during periods of high discharge (-23 to -21 ‰), and decreasing towards predominantly C3 signatures toward the end of dry periods. This suggests that high sediment mobilization during rains occurs mostly in areas with significant grassland cover (C4).