



Original article

Increased tree-growth synchronization of beech (*Fagus sylvatica* L.) in response to climate change in northwestern EuropeNicolas Latte ^{a,*}, François Lebourgeois ^{b,c}, Hugues Claessens ^a^a Forest Resources Management, University of Liège, Gembloux Agro-Bio Tech, Passage des Déportés 2, 5030 Gembloux, Belgium^b AgroParisTech, UMR1092, Laboratoire d'Étude des Ressources Forêt Bois (LERFoB), Rue Girardet 14, 54042 Nancy, France^c INRA, UMR1092, Laboratoire d'Étude des Ressources Forêt-Bois (LERFoB), Centre INRA de Nancy, F-54280, Champenoux, France

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ABSTRACT

To better understand how the radial growth of beech (*Fagus sylvatica* L.) high forests has responded to climate change, we selected 12 sites (137 trees) with optimal growing conditions along a W–E altitudinal gradient (67–590 m) in Belgium. We evaluated temporal changes in growth response to climate by using pointer year analysis, moving mean sensitivities (1860–2011), and moving bootstrapped correlation coefficients (1952–2011). The strongest driving climatic variables were identified by using the partial least squares method.

The common patterns of growth trends, pointer years, and mean sensitivities among sites provided evidences for the impact of environmental changes operating at a regional scale. The results of growth–climate analysis indicated that these changes were strongly influenced by the climatic conditions of the previous year. The climate sensitivity of beech increased progressively in response to more frequent and intense heat waves and warming-related droughts, especially during recent decades, leading to remarkable inter-site synchronization. The changes were much more pronounced for sites located in lowlands (<300 m). The differences in growth responses along the altitudinal gradient and the consequences of warming for beech growth and physiology are discussed.

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Introduction

Common beech (*Fagus sylvatica* L.) is one of the most widespread and important tree species in Europe. This species is distributed from Sicily in southern Italy to Bergen in southern Norway, and it occurs in various habitats from mountainous regions to lowlands (Fang and Lechowicz, 2006; Seynave et al., 2008). Recently, numerous tree-ring studies have revealed long-term changes in the tree-growth–climate relationships of several broadleaved and coniferous tree-species; this phenomenon is termed the “divergence problem” (D’Arrigo et al., 2008; Lebourgeois and Mérien, 2011). For beech, changes have been observed throughout Europe (Dittmar et al., 2003; Jump et al., 2006; Di Filippo et al., 2007, 2012; Friedrichs et al., 2009; Bolte et al., 2010; Scharnweber et al., 2011; Lebourgeois et al., 2012; van der Maaten, 2012; Weber et al., 2013; Castagneri et al., 2014). At high-elevation and/or high-latitude sites, temperature is the key driving factor of tree growth and observed

changes in the tree-growth–climate relationships correspond to a loss of response to temperature. At mid- and low-latitude sites, tree growth is dependent on the interaction between temperature and water availability (Lebourgeois and Mérien, 2011). Thus, to disentangle the specific warming effect of water availability on beech, it seems more appropriate to sample trees from stands with optimal growing conditions. Changes in temperature thresholds in these stands might play a major role in influencing carbon stock (Latte et al., 2013), carbon uptake projections (De Vries et al., 2006; Campioli et al., 2012) and management policies; moreover, they may induce severe economic losses (Hanewinkel et al., 2013).

Since the 1990s, anomalies in beech health (e.g., worsening crown conditions) have been observed throughout Europe (UNECE, 2005) and in Belgium (Laurent and Lecomte, 2007). During the same period, beech height growth and productivity changes have been reported in Belgium (Kint et al., 2012; Aertsen et al., 2014) and northeastern France (Bontemps et al., 2010; Charru et al., 2010). Beech is sensitive to drought during the period between budburst and the month of July and also to summer heat waves (Mund et al., 2010; Scharnweber et al., 2011). These factors are considered to be the major limiting abiotic disturbances for beech in Belgium in the

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context of climate change (Campioli et al., 2012). Since the end of the 19th century, precipitation has remained relatively stable over time; however, temperatures have increased by approximately 2 °C (Demarée et al., 2002) and will likely continue to increase in the future (Baguis et al., 2010; Collins et al., 2013). The higher frequency and intensity of heat waves and warming-related droughts have altered tree functioning and will probably continue to do so (Bréda et al., 2006; McDowell, 2011); this will be particularly true at sites with lower water availability, where beech is more sensitive to climate (Lebourgeois et al., 2005). Thus, niche-based models forecast a reduction in the beech distribution area during the 21st century, especially for lowlands in northwestern Europe (Piedallu et al., 2009; Kramer et al., 2010).

In the present study, we evaluated the influence of global warming on the tree-growth-climate relationships of northwestern European beech forests. We selected 12 beech stands along the entire altitudinal and climatic gradient of southern Belgium. To specifically investigate the effect of warming along the gradient, we defined the selection criteria so as to minimize differences between stands and growing conditions, as well as to reduce the effects of interactions between warming and water availability. Our hypotheses were that (1) global warming and related heat waves have progressively affected beech, thereby reducing its growth and increasing its sensitivity, and (2) resulting changes in growth responses to climate should vary with altitude and are probably more intense in lowlands.

Materials and methods

Study area and site and tree selection

Twelve beech stands were selected along a W–E gradient in Belgium (Fig. 1), spanning the entire elevation gradient from the lowlands (67 m) to the Ardennes Plateau (590 m). Along this gradient, the mean annual temperature decreases (from 10.3 °C to 7.2 °C) and the annual precipitation increases (from 743 mm to 1047 mm; Table 1). We restricted our stand selection to sites that were optimal for beech growth and had no noticeable topographic or soil constraints (slope $\leq 3\%$). Thus, the selected stands were among the most productive in Europe (dominant height of 19–30 m at the age of 80 years; Table 1). Soil boreholes were used to roughly estimate the maximal soil water content (≥ 90 mm; Table 1) based on the soil texture, stoniness, and depth (Ridremont et al., 2011). In addition, we limited our sampling to a fixed silvicultural context of mature and dominant trees within pure beech forests (>75% of basal area) that had been managed according to an even-aged (or regular) structure by forest administration for more than half a century. The selection criteria were thus defined to minimize differences between stands and site growing conditions, in order to focus on the influence of temperature along the altitudinal gradient. We restricted our stand selection to stands ranging in age from 84 years old to 206 years old (Table 1), in order to focus on mature trees and eliminate juvenile and senescence effects.

In each beech site, 8–15 dominant or co-dominant healthy beech trees of comparable dimensions were selected (137 trees, Table 1) and cut down in the winter of 2011–2012. One disk per tree was collected at breast height (1.3 m) at eight sites, and because of technical constraints, from higher up the stem at four sites (Table 1). The variation in sampling height did not markedly influence the results or their interpretation (Supplementary material: Figs. A and B). To facilitate handling, two bars (each 12 cm wide) were extracted from opposite side of the disk. Bars represent a good compromise between disks and cores. Bar dimensions are more practical than disks for sample preparation; moreover,

Table 1
Characteristics of the 12 beech sites. Site index = dominant height at the age of 80 years.

Site	Local site parameters		Climate parameters		Stand characteristics		Sampled tree characteristics				
	Altitude (m)	Maximum soil water content (mm)	Mean annual temperature (°C)	Annual precipitation (mm)	Mean DBH ^a (cm)	Site index (m)	Number of trees	Mean DBH ^a (cm)	DBH standard deviation (cm)	Mean cambial age (years) (min–max)	Mean sampling height (m) (min–max) ^b
STA	67	139	9.3	743	49	28.6	13	50	5	84 (70–89)	1.3
ENG	99	253	9.8	771	53	28.1	9	68	11	99 (81–123)	9 (6–12)
FLO	99	260	9.2	749	53	30.0	11	67	7	95 (89–107)	1.3
TER	112	260	10.3	774	57	26.2	13	80	8	163 (153–169)	6 (1.3–10)
LES	171	142	9.4	820	58	28.5	15	65	7	127 (118–140)	1.3
THH	225	163	9.8	827	41	24.5	13	71	7	105 (90–127)	1.3
HES	413	90	8.5	1024	41	19.3	10	62	8	131 (119–135)	1.3
POR	425	101	9.0	894	53	19.1	8	71	4	190 (153–209)	8 (4–12)
REC	458	124	8.2	994	39	20.5	12	64	6	178 (152–191)	1.3
MAR	478	105	8.0	999	38	19.3	11	70	5	206 (178–215)	6 (3–9)
SAI	521	111	8.4	971	38	20.9	10	49	8	123 (101–157)	1.3
ROC	590	98	7.2	1047	43	20.5	12	62	5	136 (111–156)	1.3

^a DBH diameter at breast height.

^b Variations in sampling height were caused by technical constraints (see text).

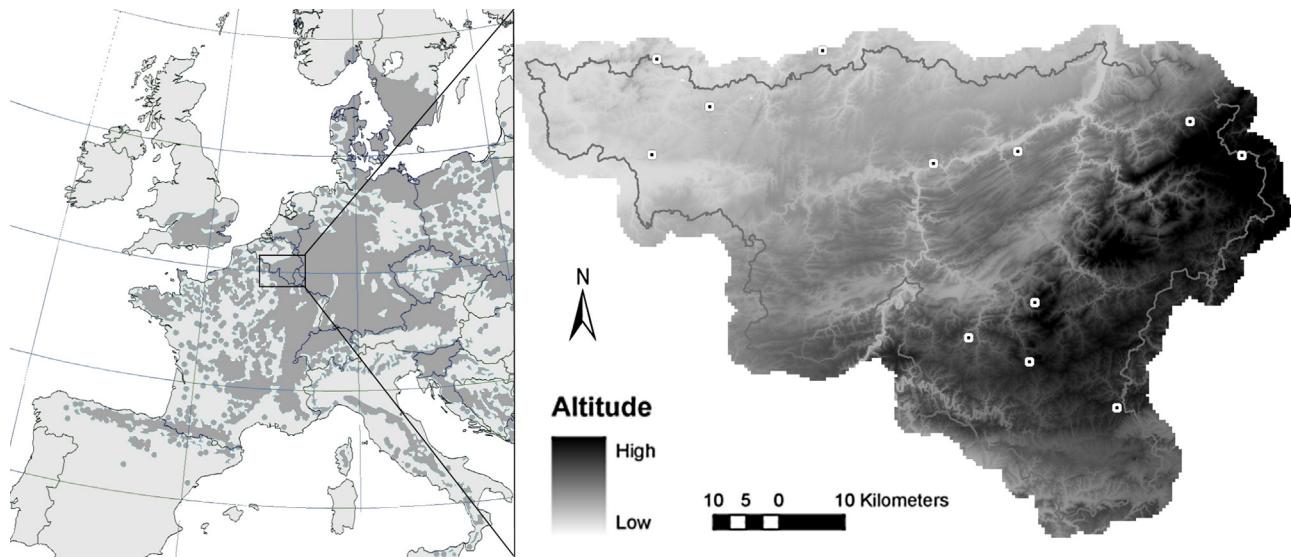


Fig. 1. Distribution map of *Fagus sylvatica* in Europe (source: www.euforgen.org) and locations of the 12 beech sites (white dots) in southern Belgium.

in comparison with cores, their higher area improves ring delineation and overall measurement quality. We ensured that the selection criteria of sites and trees (Mérian and Lebourgeois, 2011a; Mérian et al., 2013) and the resulting tree-ring series (Mérian and Lebourgeois, 2011b) met the requirements for being representative of beech growth at the scale of the study region.

Tree-ring measurement and chronology building

Air-dried bars were planed, sanded (grit 80–180), and scanned at high resolution (≥ 1200 dpi). Tree-ring widths were measured on each bar from the core to the pith with an accuracy of 1/100 mm by using WinDENDRO (Regent Instruments Canada Inc., 2009). Individual ring-width series were carefully cross-dated by progressively detecting the pointer years following Cropper's method (Cropper, 1979) as described in Neuwirth et al. (2007). A pointer year was defined as positive (or negative) when the normalized Cropper's values of at least 50% of the trees were larger (or smaller) than +1 (or -1). According to the density function of the normal distribution, the probability for a positive (or negative) growth anomaly is <16%. The cross-dating was validated for the period 1860–2011, by using available reference chronologies for Belgium (Pennincx et al., 1999; Supplementary material: Fig. C) and regional pointer years (Kint et al., 2012).

To enable more accurate quantification of wood production, ring width was converted into basal area increment (BAI) (Michelot et al., 2012; Weber et al., 2013). Site-specific chronology building and statistical calculations were conducted on BAI by using R (R Core Team, 2012) and the "dplR" package (Bunn, 2008). Individual bar series were detrended with a flexible cubic smoothing spline (frequency response of 50% at a wavelength of 24 years), to minimize low and medium frequencies caused by age, competition, and silviculture (Cook and Kairiukstis, 1990). The raw and detrended individual series were averaged by site.

Chronology statistics were computed for the common overlap period 1943–2011 (a minimum of five available trees for each site). The chronology quality was evaluated by using the expressed population signal (EPS) and the effective signal (ES). The EPS quantifies how well a chronology based on a finite number of trees represents the hypothetical perfect or true chronology, and the ES estimates the signal strength within and between trees (Wigley et al., 1984). The first-order auto-correlation coefficient

(AC) assesses the influence of the previous year's growth on the current year's growth. The mean sensitivity (MS) estimates the year-to-year variability caused primarily by climatic fluctuations. The temporal variation in site MS was calculated in a moving window of 30 years, from 1860 to 2011. The mean change in MS during the period 1943–2011 was estimated as the slope of the single linear regression of the moving 30-year MS.

Meteorological data

For each site, the daily mean, minimum, and maximum temperatures (T_{mean} , T_{min} , and T_{max} , respectively, in °C), and precipitation (P , in mm) from 1950 to 2011 were extracted from the gridded version (0.25 deg) of the European Climate Assessment Dataset (Haylock et al., 2008). The Hargreaves' potential evapotranspiration (PET) according to the formulation of Droogers and Allen (2002) and the climatic water balance ($WB = P - PET$) were computed daily. Single and multi-month climatic variables were obtained by averaging the daily values of T_{mean} , T_{min} , and T_{max} and by summing the daily values of P , PET, and WB for periods of one to seven month(s). The final month of these periods ranged from April of the previous year to October of the current year. We computed 798 variables for the period 1952–2011.

Climate-tree-growth analysis

The climate–growth relationships were analyzed for the period 1952–2011 by using R (R Core Team, 2012) in two steps. First, the best explanatory variables common to all sites were identified by using the partial least squares (PLS) regression method (R package "pls"; Mevik and Wehrens, 2007). This statistical method can be used to describe the fundamental relations between the BAI indices (Y) and the matrix of climatic variables (X), by combining the principles of principal component analysis and multiple linear regression ($Y = f(X)$). It is particularly appropriate when there are more X variables than observations (i.e., number of years) and when there is multicollinearity among X variables (Mevik and Wehrens, 2007). PLS regressions (single component) were computed between the BAI indices and standardized climatic variables (mean, 0; standard deviation, 1) for (1) all sites together, to identify the best common variables, and (2) site by site, to ensure that the identified common variables were among the best variables for each site individually.

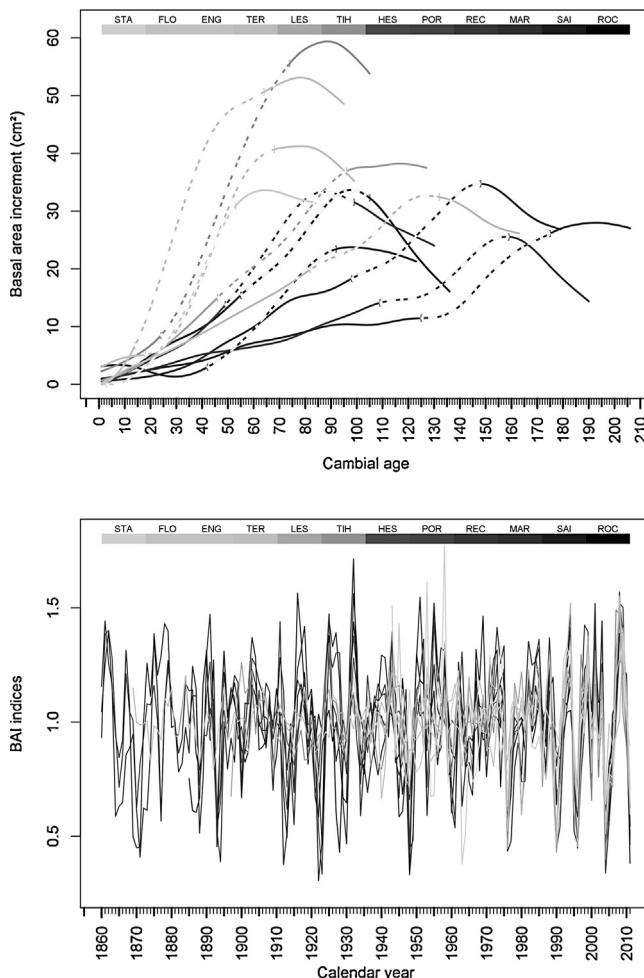


Fig. 2. Top, trends in the basal area increment (BAI) of the 12 beech sites with age (spline with a frequency response of 50% at a wavelength of 48 years). Dashed lines indicate the period 1930–1980. Bottom, BAI indices of the 12 beech sites for the period 1860–2011. Darker gray indicates higher altitudes (range, 67–590 m).

According to the variable importance in projection (Chong and Jun, 2005), one variable was selected for each data type (temperature: T_{mean} , T_{min} , and T_{max} , and water availability: P , PET, and WB). Next, the bootstrapped correlation coefficients (BCCs) between the selected variables and the site BAI index chronologies were computed by using the R package "bootRes" (Zang and Biondi, 2013) in a 30-year moving window for the period 1952–2011. The statistical significance was tested by using the bootstrap method with a 95% confidence interval (Efron and Tibshirani, 1986).

Results

Tree-ring chronologies: quality and pattern

The EPS ranged from 0.86 to 0.96 (mean = 0.93), indicating that each site chronology adequately represented the population of (co-)dominant trees (Table 2). The ES ranged from 0.40 to 0.67 (mean = 0.53), indicating that the between-trees and within-trees signals were strong. The MS values (0.14–0.30) demonstrated a relatively strong inter-annual variability within tree-ring indices. The AC values (0.41–0.87) revealed a high dependency of the current year's growth on the previous year's growth. The BAIs of the 12 beech sites were in proportion to the site index (Table 1) and showed a similar pattern of change over time (Fig. 2, top). The

growth rate accelerated between the 1930s and the 1970s, and thereafter, strongly decelerated.

Pointer year analysis

For the common period 1943–2011, the number of positive and negative pointer years (PYs) ranged from 5 to 20 according to site; moreover, negative PYs occurred more frequently than positive PYs (Table 2). From 1860 to the beginning of the 1990s (Fig. 3), the proportion of sites with common PYs was <75% (except in 1976). During the two last decades, a much higher frequency of PYs (particularly negative PYs) was observed; 58% of PYs occurring from 1943 to 2011 were concentrated during the period 1990–2011. Six negative PYs were common to at least eight sites as follows: 1948 (9), 1976 (11), 1990 (8), 1996 (12), 2004 (11), and 2011 (11). There were many points of convergence between these major PYs and extreme climatic events in Belgium (Institut Royal Météorologique (IRM), 2000). The main growth losses were observed during the current year if drought occurred in spring (e.g., 1990, 2011) or during the following year if a heat wave (and drought) occurred in summer (e.g., 1948, 2004). Exceptional growth losses were induced by the combined effects of stressful growing conditions during the previous summer and current spring (1976, 1996). Positive years were generally related to favorable conditions in the previous summer (wet and/or cool).

Mean sensitivity increase and growth synchronization

For the period 1860–2011, the mean sensitivity (MS) varied according to the site altitude (Fig. 4). Before the 1970s, the MS of lower altitude (LA, <300 m) sites was lower than that of higher altitude (HA, >400 m) sites; this phenomenon occurred despite the fact that LA sites had lower precipitation and warmer temperatures but higher soil water contents than did the HA sites (Table 1). After the 1970s, the MS of LA sites (with the exception of STA) increased more intensely to reach the same level as that of the HA sites.

For the common period 1943–2011, the MS was higher for older stands at LA and HA sites; however, the MS changes differed between the LA and HA sites (Fig. 4). The MS increases were higher for older stands at LA sites and for younger stands at HA sites. During the same period, the 12 BAI index chronologies (Fig. 2, bottom) became increasingly synchronized, leading to a strong increase in the inter-site correlation for the two altitude groups (Fig. 5).

Climate–tree-growth relationships

For the period 1952–2011, the r -square of the "all sites" PLS was 21%, and the r -square values of the "site by site" PLSs ranged from 17% to 39% and were systematically higher for the HA sites (Table 2). The main driving climatic variables were the same for the "all sites" PLS and the mean of the 12 "site by site" PLSs and were (1) the mean of the daily maximum temperature from July to October of the previous year (TMAX), and (2) the sum of the daily climatic water balance from May to November of the previous year (CWB). The remaining variables did not explain a significant part of the common inter-site growth variation.

The bootstrapped correlation coefficient (BCC) between the 12 BAI index chronologies and the best explanatory variables was significant for the period 1952–2011; moreover, pronounced changes were observed over time (Fig. 6). For TMAX, all sites showed a similar pattern, with a strong increase in the absolute BCC toward the end of the 1970s, followed by high and significant values. From 1952 to 2011, the mean TMAX (all sites) increased continually by +0.8 °C (+1.4 °C in July and August). The 10–90th percentile range of TMAX also increased, but to a much larger extent (up to +40%) at LA

Table 2

Statistics of the raw and detrended chronologies of the 12 beech sites for the common period 1943–2011.

Site	Non-detrended chronologies			Number of pointer years		Detrended chronologies			R-square PLS (%)
	Averaged BAI (cm ² /year)	Averaged ring-width (mm/year)	First-order auto-correlation coefficient (AC)	Negative	Positive	Mean sensitivity (MS)	Effective signal (ES)	Expressed population signal (EPS)	
STA	22.9	2.7	0.87	6	5	0.24	0.54	0.94	17
ENG	46.5	3.0	0.58	3	2	0.14	0.41	0.86	17
FLO	33.2	3.2	0.81	7	5	0.14	0.40	0.89	18
TER	28.6	2.1	0.41	10	7	0.24	0.53	0.94	31
LES	32.5	2.3	0.54	9	5	0.23	0.51	0.94	39
TIH	47.8	3.0	0.57	8	5	0.21	0.56	0.94	32
HES	28.8	2.2	0.49	10	10	0.24	0.66	0.95	27
POR	20.1	1.6	0.56	11	7	0.28	0.54	0.92	33
REC	29.1	1.7	0.42	10	4	0.29	0.54	0.94	38
MAR	23.5	1.5	0.59	11	7	0.30	0.67	0.96	39
SAI	19.2	1.7	0.73	5	5	0.19	0.51	0.91	34
ROC	26.3	2.0	0.60	8	6	0.23	0.53	0.93	29
All sites	29.9	2.2	0.60	8.2	5.7	0.23	0.53	0.93	29
Lower elevation sites (<300 m)	35.2	2.7	0.63	7.2	4.8	0.20	0.49	0.92	26
Higher elevation sites (>400 m)	24.5	1.8	0.56	9.2	6.5	0.26	0.58	0.94	33

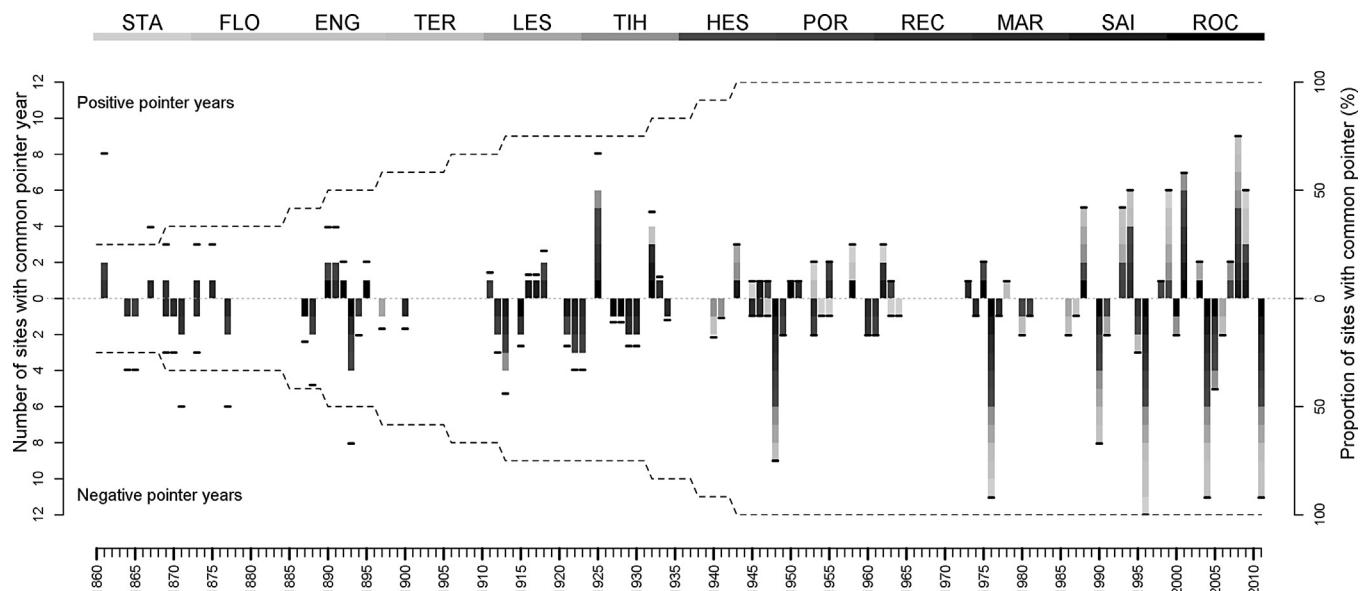


Fig. 3. Number (left y-axis, bar plot) and proportion (right y-axis, hyphens) of sites with common positive and negative pointer years among the 12 beech sites for the period 1860–2011. The bar height corresponds to the number of sites; the shading within each bar indicates site and altitude; darker gray indicates higher altitudes (range, 67–590 m). The dotted lines correspond to the number of sites (sample depth) with at least five trees available by site.

sites. For CWB, the mean value remained relatively constant over time; however, the BCCs increased for all sites. The three youngest LA sites (FLO, STA, and ENG) showed significant BCC values later than did the other investigated sites (Fig. 6) and this confirmed the smaller MS increase for young stands at LA sites (Fig. 4).

Discussion

In the present study, the common patterns of basal area increment (Fig. 2), negative pointer years (Fig. 3), mean sensitivity (Fig. 4), inter-site correlations (Fig. 5), and bootstrapped correlations (Fig. 6) along the altitudinal gradient provide evidence for the impact of environmental changes on beech at a regional scale.

The overall growth increase observed at the European scale before the 1970s has frequently been attributed to increasing temperatures combined with nitrogen deposition and increasing atmospheric CO₂, as well as to changes in forest management

(e.g., De Vries et al., 2006; Bontemps et al., 2009, 2011). The stand selection was oriented toward reducing the impact of forest management as much as possible. However, historical documents mention that before the 1950s, beech stands were generally managed at high density (e.g., a 24-year or less frequent thinning cycle) and that subsequent forest administration intensified the use of silvicultural practices (e.g., a 12-year thinning cycle). Previous studies reported that thinning can enhance the growth and resilience of mature beech trees (Le Goff and Ottorini, 1999; van der Maaten, 2013). Therefore, changes in forest management after the 1950s may partly explain the observed growth increase before the 1970s.

After the 1970s, the observed growth decrease is consistent with the results of recent studies on beech productivity in northern Belgium (Kint et al., 2012; Aertsen et al., 2014) and northeastern France (Bontemps et al., 2010; Charru et al., 2010). Tree-ring studies of beech in different regions of Europe also reported decreasing radial growth and increasing climate sensitivity (Dittmar et al.,

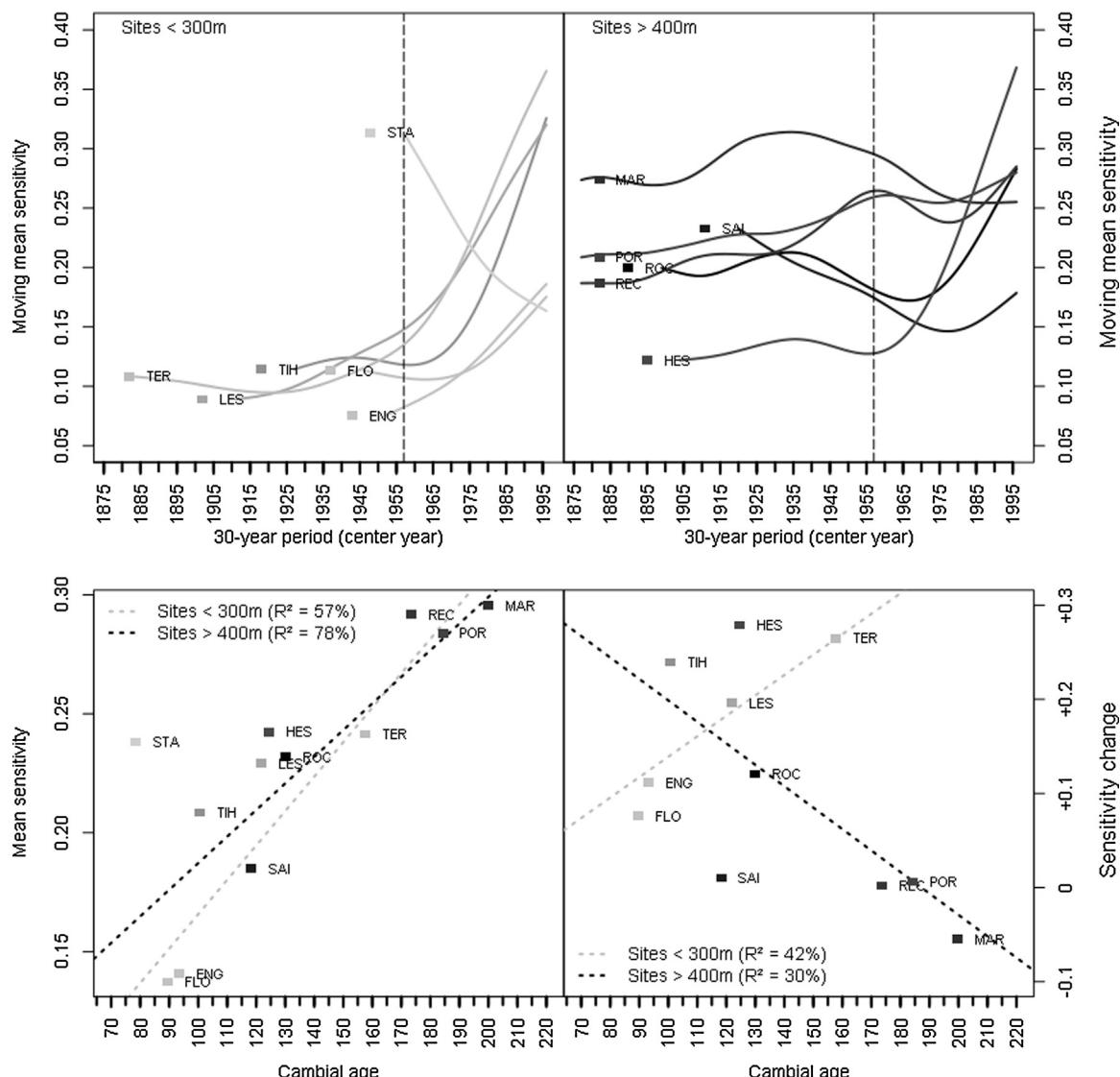


Fig. 4. Top, trends in the moving mean sensitivity (MS) of the 12 BAI indices performed in a 30-year moving window for the period 1860–2011 (spline with a frequency response of 50% at a wavelength of 48 years). Vertical dotted lines indicate the beginning of the common period 1943–2011. Bottom, scatter plot of site-MS with cambial age (left), and its change over time (right) for the period 1943–2011. Dotted lines correspond to single linear regressions excluding the STA site.

2003; Jump et al., 2006; Di Filippo et al., 2007, 2012; Friedrichs et al., 2009; Scharnweber et al., 2011; Lebourgeois et al., 2012; van der Maaten, 2012; Weber et al., 2013; Castagneri et al., 2014), except in high mountains (Di Filippo et al., 2012) and at high latitudes (Bolte et al., 2010). However, the rapidly decreasing growth that occurred uniformly along the climatic gradient was unexpected, because the stands were selected at sites that were optimal for beech growth, had no noticeable topographic and soil constraints, and were located in an area that was distant from the distribution limit; these stands were among the most productive in Europe.

On the basis of the continuity after the 1950s, the growth decline cannot be explained by forest management. The aging of trees is frequently related to a decline in forest productivity and is characterized by an increase in sensitivity to perturbations (Genet et al., 2009). The results of our present study confirm the increasing sensitivity of beech with age at the regional scale; however, the influence of aging on sensitivity changes over time was inconsistent along the altitudinal gradient. On average, the sensitivity of older stands (>160 years old) remained stable, whereas it increased for the remaining stands. Our findings indicate that the

influence of aging is limited and cannot solely explain the observed changes. This assumption is reinforced by the fact that the sensitivity increase was triggered in a large range of tree ages (70–215 years) during a relatively short period. Recently, Kint et al. (2012) reported a negative effect of excess nitrogen deposition on beech productivity in northern Belgium; however, climate variables (particularly decreasing air humidity and increasing temperature) were essential for modeling the growth decrease. Furthermore, growth declines were observed in other regions of northwestern Europe (Bontemps et al., 2010; Charru et al., 2010), where the rate of nitrogen deposition was lower than in Belgium (De Vries et al., 2003).

The results of correlation analysis clearly revealed the increasing influence of the previous year's climatic conditions (Dittmar et al., 2003; Lebourgeois et al., 2005; Friedrichs et al., 2009; Härdtle et al., 2013). Several ecophysiological hypotheses can be used to explain the negative impact of warmer temperatures during summer and at the start of autumn. (1) Increasingly high temperatures exacerbate the effects of heat waves and droughts on beech growth (Lebourgeois et al., 2005) and vitality (Bréda et al., 2006), and therefore increase the risk of hydraulic failure (McDowell, 2011). Beech

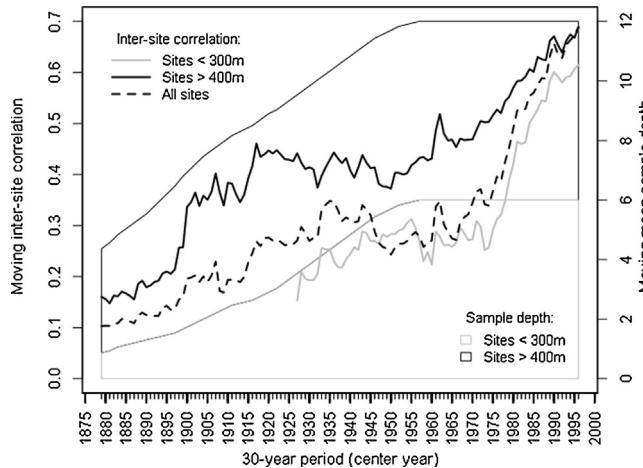


Fig. 5. Left y-axis, inter-site correlation coefficient between the BAI indices of at least two sites at lower altitudes (<300 m, in gray) and higher altitudes (>400 m, in black) performed in a 30-year moving window for the period 1860–2011. Right y-axis, mean number of sites available (with at least five trees) for the considered year.

defoliation is particularly high in a year following a warm summer, and leaf area is frequently lower for several years following a severe drought (Seidling, 2007). (2) Warmer autumnal temperatures extend beech canopy duration (Vitasse et al., 2009) and thus

may limit the formation of metabolic reserves (because respiration exceeds photosynthesis) and consequently growth in the following year (Michelot et al., 2012). (3) Fine roots of beech trees have a very high turnover rate in dry years (Meier and Leuschner, 2008). Thus, a higher frequency of warming-related droughts may reduce the amount of carbohydrates available for growth and to restore fine root stock for water uptake (Genet et al., 2009). (4) High temperatures (Drobyshev et al., 2010; Mund et al., 2010) and drought (Piovesan and Adams, 2001) in the summer can trigger beech masting in the following year. Thus, the observed higher frequency of mastings (data not published) may accentuate the carbon depletion.

In the present study, the mean and variability of high temperatures in summer and autumn increased over time, whereas the water balance during the growing season remained relatively stable. Thus, increasing sensitivity was most likely driven by the higher frequency and intensity of heat waves and warming-related droughts. The results of pointer year analysis confirmed that the more frequent growth losses, which occurred uniformly at the regional scale, led to a pronounced inter-site synchronization. After the 1990s, the synchronization was so high that the site chronologies could be interpreted as tree chronologies of a unique stand. This trend may have been accentuated by tropospheric ozone (Dittmar et al., 2003; Matyssek et al., 2010). Concentrations of tropospheric ozone are strongly correlated with high temperatures (Stathopoulou et al., 2008) and have increased in Belgium because of increasing air pollution (De Vries et al., 2003). The

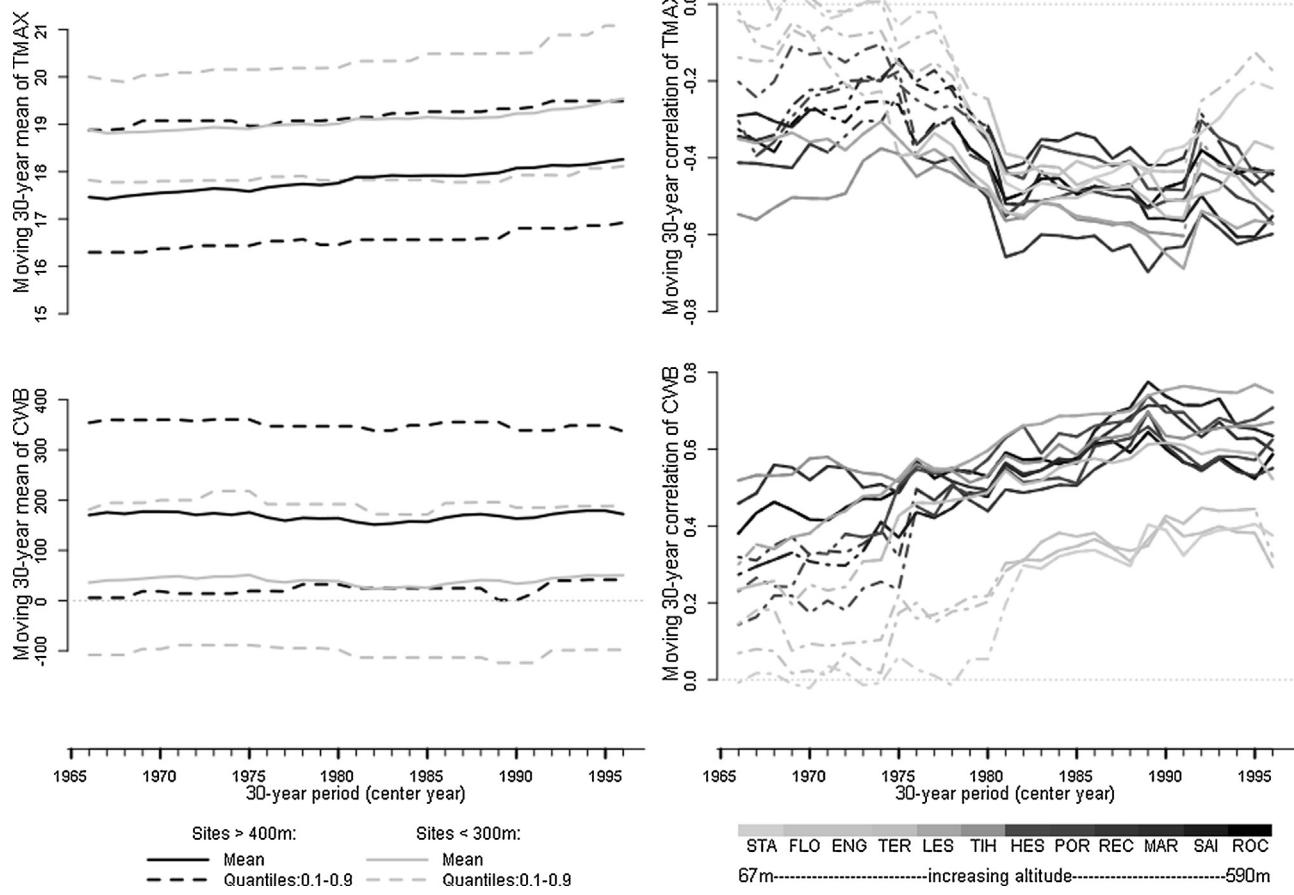


Fig. 6. Left, Moving 10th–90th percentiles of the means of the two climatic variables (TMAX, CWB) for lower altitude sites (<300 m, in gray) and higher altitude sites (>400 m, in black) performed in a 30-year moving window for the period 1952–2011. TMAX is the mean of the daily maximum temperature from July to October of the previous year. CWB is the sum of the daily climatic water balance from May to November of the previous year. Right, moving bootstrapped correlation coefficients between the 12 site BAI indices and the two climatic variables (TMAX, CWB) performed in a 30-year moving window for the period 1952–2011. Darker gray indicates higher altitudes (range, 67–590 m). Solid lines indicate a significant correlation with a 95% confidence interval.

later and higher synchronization observed in lowlands may have been induced by higher temperature variability and/or differences in water availability. The higher maximal soil water content (but lower precipitation) at lower altitudes may have ceased to mitigate the warming effects later than at higher altitudes. Our results are in accordance with those of Friedrichs et al. (2009) and Weber et al. (2013), who reported that recent changes in beech sensitivity were more pronounced at sites with higher soil moisture.

The overall trend of decreasing growth observed during recent decades was mainly derived from the higher frequency of negative pointer years; on average, the intense growth losses were not fully compensated by the good growth rate in favorable years. However, in accordance with the observations of van der Werf et al. (2007), growth recovered in the following years, implying that, to date, beech has tolerated its changing environment. The observed changes in beech growth and response to climate at sites with optimal growing conditions indicate that the positive influence of local growing conditions (particularly water availability) decreased with warming. Given the predicted warming on the pan-European scale, beech is likely to be affected throughout a large part of its distribution area (excluding high elevations and high latitudes).

Conclusions

In the present study, we used a dendroecological approach to investigate the influence of climate change on beech growth under optimal growing conditions (high soil fertility and high water availability) along an altitudinal gradient at a regional scale (southern Belgium). We observed common patterns of decreasing growth and pronounced increasing climate sensitivity and inter-site synchronization since the 1980–1990s. Our results provide consistent evidence that changes in the tree-growth–climate relationships are induced by more frequent and intense summer heat waves and warming-related droughts. We found that beech is becoming more strongly affected than in the past, particularly in the lowlands, where the phenomena is concomitant with a decrease in tree vitality (Laurent and Lecomte, 2007; Supplementary material: Fig. D).

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Appendix A. Supplementary data

Supplementary material related to this article can be found, in the online version, at <http://dx.doi.org/10.1016/j.dendro.2015.01.002>.

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