

**1 Ten years of Martian nitric oxide nightglow  
2 observations**

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3 We present ten years of Martian NO nightglow SPICAM observations in  
4 limb and stellar occultation modes.

5 The NO nightglow is used as a tracer of the summer-to-winter hemispher-  
6 ical circulation in the upper atmosphere of Mars. Its distribution roughly fol-  
7 lows the curve  $\text{latitude} = -80 \sin(\text{solar longitude})$ , with deviations. We find  
8 that the peak brightness is  $5 \pm 4.5$  kR, situated at  $72 \pm 10.4$  km. It ranges  
9 from 0.23 to 18.51 kR and from 42 to 97 km. These values are consistent with  
10 previous studies. We also present maps of the brightness of the NO emission  
11 peak and its variability, an important factor that can reach up to 50% of the  
12 emission and is not reproduced by average brightness model maps. The char-  
13 acteristics and factors that may control the emission are investigated. In par-  
14 ticular, we show that the solar activity exerts a positive influence on the num-  
15 ber of detections. It does not influence, on the contrary, the brightness or  
16 altitude of the peak of the NO nightglow emission.

17 Results presented in this study lead to future comparisons with global Mar-  
18 tian atmospheric models and observational targets for the IUVS-MAVEN.

## 1. Introduction

19 The upper atmosphere of Mars dynamics, energy balance, structure and composition  
20 depend on its multiple interactions with the lower atmosphere and the ionosphere. Its  
21 study enhances our understanding of the atmosphere and its coupling with the solar  
22 forcing. The upper atmosphere is the major target of present and future Martian  
23 missions as the NASA Mars Atmosphere and Volatile Evolution (MAVEN) spacecraft.  
24 The study of planetary airglow provides valuable information concerning the atmosphere  
25 where it is produced as these emission remotely probe the composition, temperature and  
26 dynamics of an atmosphere.

27 In the dayside thermosphere of Mars, the extreme ultraviolet solar radiations photodis-  
28 sociate  $\text{CO}_2$  and  $\text{N}_2$  molecules.  $\text{O}(^3\text{P})$  and  $\text{N}(^4\text{S})$  atoms are then carried by the summer-to-  
29 winter hemispheric transport. They recombine to form  $\text{NO}(\text{C}^2\Pi)$  excited molecules that  
30 directly emit the UV  $\delta$  and  $\gamma$  bands (the  $\delta$  bands are emissions of the  $\text{C}^2\Pi$  state, while  
31 the  $\gamma$  bands are emissions of the  $\text{A}^2\Sigma$  state, which has been populated by cascading from  
32 the  $\text{C}^2\Pi$  state): these emissions are indicators of the N and O atom fluxes transported by  
33 the summer-to-winter dayside to nightside Hadley cell.

34 The first detection of the nitric oxide UV airglow on Mars nightside was reported by  
35 Bertaux et al. (2005) using the SPICAM (Spectroscopy for Investigation of Characteristics  
36 of the Atmosphere of Mars) spectrograph on board Mars Express (MEx). They observed  
37 an emission peak reaching 2.2 kR located at 70 km. The limiting factor for this emission is  
38 the nitrogen atom flux descending towards the atmospheric layer where N atoms recombine  
39 with O to produce  $\text{NO}^*$ . They estimated this downward flux to be  $2.5 \times 10^8$  atoms  $\text{cm}^{-2}$

40  $s^{-1}$ , about a third of the production of N atoms by EUV photodissociation of  $N_2$  molecules  
41 on the dayside.

42 Cox et al. (2008) looked for correlations between the emission peak brightness and  
43 altitude with several factors that may affect the emission rates, such as: latitude, local  
44 time, magnetic field and solar activity. They noticed that none of these factors seems to  
45 control the emission, which exhibit large variations. The dataset used by Cox et al. (2008)  
46 included 21 airglow detections between August 2004 and May 2006. The characteristics  
47 (brightness and altitude) of the NO emission peak from the study of Cox et al. (2008)  
48 are summarized in Table 1. Cox et al. (2008) compared observational emission profiles  
49 with the results of a one-dimensional chemical-diffusive model in which the continuity  
50 equations for  $O(^3P)$  and  $N(^4S)$  and NO are used to determine the eddy diffusion, oxygen  
51 and nitrogen density profiles and the vertical downward nitrogen flux.

52 Gagné et al. (2013) used 2275 SPICAM stellar occultations accumulated between June  
53 2004 and September 2009 to analyze 128 detections of the NO nightglow. They noticed  
54 an interannual variability of the number of detection of the emission, linked to changes  
55 in the solar flux during that time period . The number of detections increases with the  
56 solar flux, in agreement with the paradigm of production of  $N(^4S)$  on the dayside. They  
57 analyzed the peak intensity and altitude of the NO emission see Table 1. They explained  
58 that the mean brightness they observed is higher than the value found by Cox et al. (2008)  
59 as the result of two factors: the dataset they used is larger than the dataset of Cox et al.  
60 (2008) and it contains observations in various seasons covering three Martian years. The  
61 observations they analyzed were obtained in large part at higher solar activity. Gagné

62 et al. (2013) also noticed that the peak altitude is statistically lower in the southern  
63 hemisphere. This hemispheric asymmetry was not reproduced by the LMD (Laboratoire  
64 de Météorologie Dynamique) model described by Gonzalèz-Galindo et al. (2009) and  
65 Lopez-Valverde et al. (2011). No correlation was found between the altitude and the  
66 brightness of the peak. They explained that this is caused by the fact that the emission is  
67 localized in regions where downward fluxes of N and O atoms are important. In agreement  
68 with the LMD results, the detections of the NO  $\delta$  and  $\gamma$  bands are roughly located along  
69 the curve latitude =  $-80 \sin(L_s)$ , with outliers (detections away from the curve) and  
70 non-detections along the curve. The LMD model also predicts a brightness at the winter  
71 poles exceeding 100 kR, which was never detected. Finally, Gagné et al. (2013) pointed  
72 out another discrepancy between the data and the model: the large variability for the  
73 altitude of the peak is not reproduced by the LMD model.

74 We here use detections and non-detections of the NO  $\delta$  and  $\gamma$  bands by SPICAM to  
75 investigate the dynamics of the nightside upper atmosphere of Mars. The dataset used  
76 in this study covers the years 2003-2013, almost a full solar cycle. Results shown in  
77 this study will provide comparative information useful for the future observations of the  
78 Martian UV nightglow by the Imaging UltraViolet Spectrograph (IUVS) on board the  
79 MAVEN spacecraft and useful information for future improvements of GCMs.

## 2. Observations

80 The Mars Express spacecraft travels along a nearly polar eccentric orbit with a period  
81 of 6.72 hour, a periapsis of about 300 km and an apoapsis of 10,100 km. The SPICAM  
82 instrument on board Mars Express is composed of both an UV and an infrared spectrom-

83 eter. The UV spectrometer covers the range from 118 to 320 nm, which includes the  
84 totality of the NO  $\delta$  and  $\gamma$  bands, from 190 to 300 nm.

85 We use measurements from the UV spectrometer SPICAM in limb profile mode, de-  
86 scribed by Bertaux et al. (2006) and Cox et al. (2008). A typical observation lasts  $\sim$ 20  
87 minutes with one spectrum recorded every second in each of the 5 spatial bins (adjacent  
88 segments of the CCD) of the instrument. A spectrum can be collected after photons travel  
89 through either a small (50  $\mu\text{m}$ ) or a wide (500  $\mu\text{m}$ ) slit, providing a spectral resolution of  
90 1.5 and 6 nm respectively. The spatial vertical resolution depends on the distance between  
91 the spacecraft and the atmosphere of Mars, and may be as small as a few kilometers when  
92 the spacecraft is close to the planet. The field of view of a single SPICAM pixel is 40x40  
93 arcsec.

94 This study also includes the SPICAM observations performed in stellar occultation  
95 mode, as described by Bertaux et al. (2006) and Gagné et al. (2013). During a stellar  
96 occultation observation, the spacecraft pointer is directed to a star, hence providing an  
97 absolute calibration of the emission by subtracting the known star brightness. The tech-  
98 nique used to retrieve the NO emissions from stellar occultations was described by Royer  
99 et al. (2010) and developed to study the nitric oxide emissions in the upper nightside  
100 atmosphere of Venus.

101 These two techniques provide a large dataset of 5000 observations among which more  
102 than 200 present identifiable NO emissions. In the limb viewing mode, 700 observations  
103 are contaminated by photons from the bright dayside, which is orders of magnitude  
104 brighter than the nitric oxide emission. These 700 observations have therefore been ignored

105 in the data analysis. 111 observations do not allow defining the peak brightness and  
106 altitude because of very low emissions, which are typically under 0.2 kR. 1789 observations  
107 show no detectable NO emission.

### 3. Results

108 We here analyze correlations between the peak brightness and altitude and factors  
109 such as geographical location and solar flux influence. A large inhomogeneity appears in  
110 the detection of the nitric oxide  $\delta$  and  $\gamma$  bands. Figure 1 shows the distribution of the  
111 observations (panel a) and ratio of the number of positive detections by the number of  
112 observations (panel b) performed by SPICAM in the limb viewing mode for different F10.7  
113 indices encountered during the 2003-2013 period. We present two sets of F10.7 values.  
114 The upper axis shows the value of the F10.7 index recorded at the time of the observation  
115 at Earth. The lower axis shows the F10.7 solar flux corrected for the distance between  
116 the Sun and Mars and the solar longitude of Mars during the observations with respect to  
117 the solar longitude of the Earth. The latter values also take into account the eccentricity  
118 of the Martian orbit. Gagné et al. (2013) showed that the number of detections increases  
119 with the solar flux. This result was reproduced by the LMD simulations. We statistically  
120 confirm this long term variability. Figure 1 panel a shows a decrease in the number of  
121 observations as the solar flux increases. We show in Figure 1b the ratio between the  
122 number of positive detections of the NO airglow and the total number of detections, for  
123 increasing solar fluxes. This is an indicator of more numerous positive detections for higher  
124 solar activity conditions. No relation between the solar activity and the peak brightness

125 was however found. The analysis of this unexpected result is beyond the scope of this  
126 paper and is a potential topic for future work.

127 Cox et al. (2008) showed the lack of correlation between the peak altitude and bright-  
128 ness. This result was then confirmed by Gagné et al. (2013). Cox et al. (2008) found a  
129 mean peak brightness and altitude of the NO emission of  $1.2 \pm 1.5$  kR at  $73 \pm 8.2$  km.  
130 Gagné et al. (2013) found  $4 \pm 3.5$  kR at  $83 \pm 24$  km. Here, we find an average peak  
131 for the nitric oxide  $\delta$  and  $\gamma$  bands of  $5 \pm 4.5$  kR located at  $72 \pm 10.4$  km (see Table 1).  
132 We find identical results when using the same dataset as Cox et al. (2008). Our results  
133 are consistent with those from the study led by Gagné et al. (2013). The altitude of the  
134 peak is consistent in the three studies. The peak brightness and altitudes range from 0.23  
135 to 18.51 kR and from 42 to 97 km. The emission layer (from 40 to 100 km) is larger  
136 than the one (from 60 to 80 km) predicted by the LMD model (González-Galindo et al.  
137 (2009)), but in good agreement with the results obtained from stellar occultations. One  
138 preliminary result found by Cox et al. (2008) is not reproduced in this extensive study:  
139 the altitude of the emission peak does not seem to be controlled by the planetocentric lat-  
140 itudinal position of the emission (Stiepen, 2014). Finally, peaks in the emission profiles  
141 close to the winter poles are brighter than those near the equator.

142 Both SPICAM stellar occultations and limb viewing observations are represented in a  
143 latitude/solar longitude map in Figure 2. In panel 2a, triangles represent the detections  
144 in stellar occultation mode while diamonds are detections in limb viewing mode. In panel  
145 2b, grey dots show the locations of the non-detections in stellar occultation mode and  
146 black dots refer to non-detections in limb viewing mode.



147 The analysis of the brightness at different latitudes and seasons is presented in Figure  
148 3. In figure 3, all observations are combined to construct an extensive dataset of  $\sim 200$   
149 detections. Panel a shows the mean brightness in each  $5^\circ$  latitude/  $10^\circ$  solar longitude  
150 bin, panel b shows the  $1\text{-}\sigma$  variability of the brightness in each bin, and panel c shows the  
151 number of observations in each bin.

152 Figures 2 and 3 summarize all observations and compare with the outputs from the  
153 LMD model (see figure 5 from Gagné et al., 2013). The model roughly reproduces the  
154 location where the NO nightglow is detected. There are however many non-detections  
155 within regions where the NO airglow is predicted to be bright and detections have been  
156 made in the equatorial region. Figure 3 shows an analysis of the variability of the emission.  
157 The LMD model reproduces well the mean brightness of the emission (Figure 3, panel  
158 a). The mean number of observations in each bin is  $\sim 10$ , as shown in Figure 3 panel c.  
159 The standard deviation of the mean brightness illustrates the variability of the emission  
160 for similar conditions (Figure 3 panel b). The variability can reach 8 kR, slightly less  
161 than 50% of the peak brightness of the brightest profiles. This variability cannot be  
162 reproduced in averaged simulations. The causes of this variability is unknown. Potential  
163 candidates to explain this variability include Eddy diffusion, wave drag and changes in  
164 the global circulation. This variability is an important constraint for future developments  
165 of atmospheric models.

#### 4. Conclusions

166 The SPICAM instrument on board Mars Express spacecraft has observed the nitric  
167 oxide emission in two different viewing modes: tangential limb and stellar occultations.

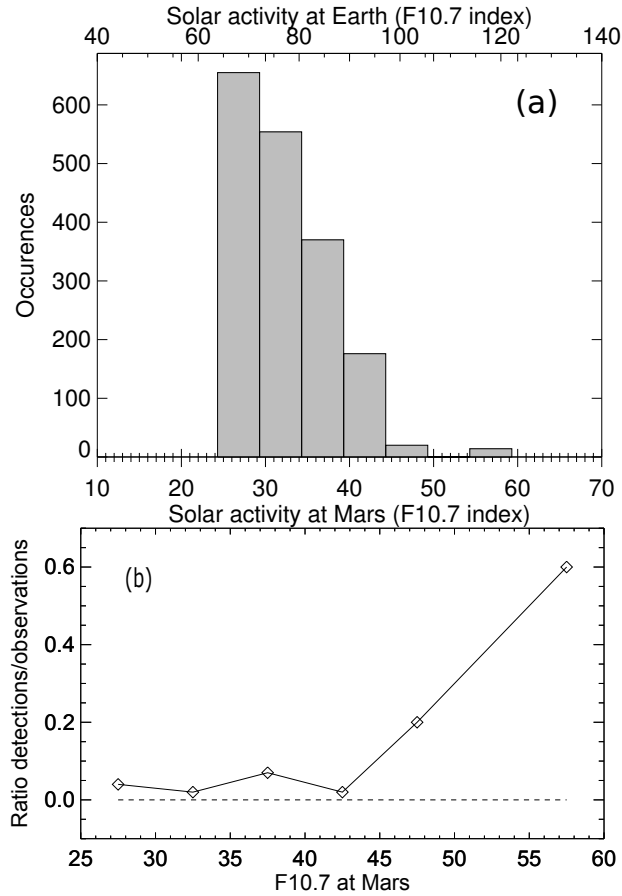
168 We merged the two datasets to carry out an extensive survey of the nitric oxide  $\delta$  and  $\gamma$   
169 bands nightglow for almost a full solar cycle (2003-2013). The nitric oxide nightglow is a  
170 tracer of the dynamics of the upper atmosphere of Mars, useful to constrain the summer-  
171 to-winter hemispherical transport. We show that the number of detections increases with  
172 the solar activity, despite a lack of correlation between the F10.7 index and the brightness  
173 of the NO nightglow. The detection rate is also higher in specific regions: they are more  
174 frequent closer to the winter pole. The detections seem to be roughly distributed along a  
175 latitude =  $-80 \sin(L_s)$  curve, in agreement with LMD simulations. The characteristics of  
176 the emission peak are analyzed and no correlation with geographical or solar flux related  
177 factors is found, in contradiction with conclusions from Cox et al. (2008). We find the  
178 peak of the nitric oxide  $\delta$  and  $\gamma$  bands vertical profiles of  $5 \pm 4.5$  kR situated at  $72 \pm 10.4$   
179 km. The peak brightness and altitudes range from 0.23 to 18.51 kR and from 42 to 97 km.  
180 We have constructed maps of the brightness of the nitric oxide emission and its variability,  
181 which show that the emission is highly variable, even for similar conditions, with variations  
182 that may reach 50% of the brightest profiles. This is an indicator of variations in  
183 the N fluxes at time scales shorter than a Martian year. Similarly, discrepancies in the  
184 regions of the detections between the data and the model are indicators of short-term  
185 variations of the N flux or the circulation pattern likely caused by changes in the Eddy  
186 diffusion, the wave drag and the global circulation. These questions will be investigated  
187 in future comparisons between the data and the model. This study may also define future  
188 investigations using the IUVS-MAVEN measurements of the nitric oxide nightglow.

189 **Acknowledgments.** The authors thank the SPICAM and the Mars Express teams  
190 for the excellent quality of their work. A. Stiepen was supported by the PRODEX  
191 program of the European Space Agency (ESA) managed with the help of the Belgian  
192 Space Policy Office (BELSPO), Belgian American Education Foundation and Rotary  
193 District 1630. Data used in this study are available from ESA planetary science archives  
194 ([http://www.rssd.esa.int/index.php?project=MARSEXPRESS&page=file\\_tracker&reversed=on](http://www.rssd.esa.int/index.php?project=MARSEXPRESS&page=file_tracker&reversed=on)).  
195 The authors would also like to thank Jane Fox for her useful comments.

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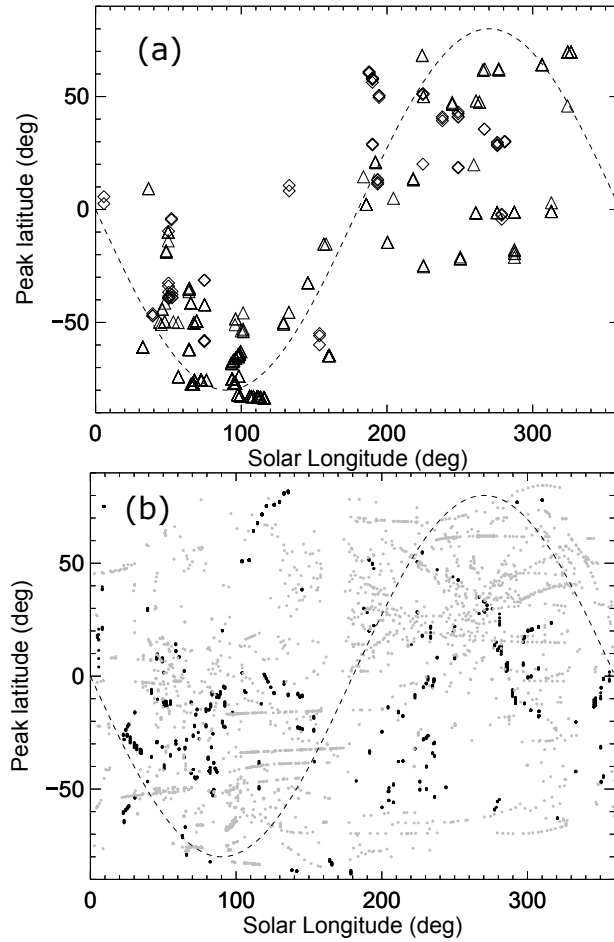
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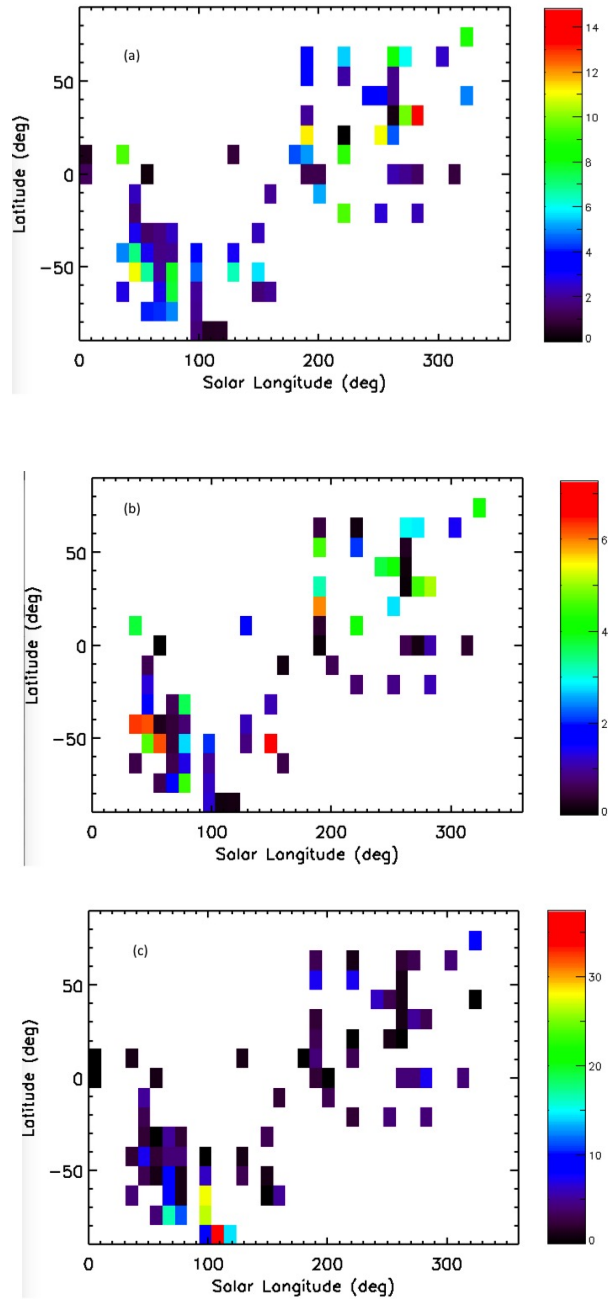
**Figure 1.** Distribution of the NO  $\delta$  and  $\gamma$  bands detections as a function of solar activity. The distribution SPICAM observations is shown on panel a. Panel b shows the ratio between the positive detections of the NO emission and all SPICAM observations. Solar activity is represented by the F10.7 index values at Earth and at Mars.

**Table 1.** NO nightglow peak characteristics

	Cox et al. (2008)	Gagné et al. (2013)	This study
Mean peak brightness (kR)	1.2	4	5
Standard deviation (kR)	1.5	3.5	4.5
Peak brightness range (kR)	0.2 - 10.5	0.5 - 10	0.23 - 18.51
Mean peak altitude (km)	73	83	72
Standard deviation (km)	8.2	24	10.4
Peak altitude range (km)	55 - 92	40 - 130	42 - 97



**Figure 2.** Mapping of the nitric oxide detections. A latitude/solar longitude map of the NO observations performed in limb scan and stellar occultation modes is presented. Panel a shows the detections of the NO nightglow. Triangles are detections performed in stellar occultation mode and diamonds show the detection made in limb viewing mode. Panel b shows the non-detections. Grey dots refer to the stellar occultation mode and black dots show the non-detections in limb viewing mode. In both panels, the line represents the curve  $\text{latitude} = -80 \sin(L_s)$ . Less than 5% of the observations led to detection of NO nightglow.



**Figure 3.** Mapping of the nitric oxide intensity and variability. Panel a shows the brightness of the peak averaged in  $5^\circ$  latitude and  $10^\circ$  solar longitude bins. The one-sigma standard deviation of these mean values is showed in panel b. In panel a and panel b, the color bar indicates the intensity in kR. Panel c shows the number of observations within each bin. In this panel, the color code indicates the number of occurrences. All panels

include NO nightglow observations from limb viewing and stellar occultations modes.

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January 15, 2015, 12:19pm

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