

# Long-term trends of NO<sub>y</sub> above northern mid-latitudes as inferred from Jungfraujoch, HALOE and ACE-FTS solar observations.



## INTRODUCTION

The NO<sub>y</sub> family of gases, defined as NO + NO<sub>2</sub> + NO<sub>3</sub> + 2×N<sub>2</sub>O<sub>5</sub> + HNO<sub>3</sub> + HNO<sub>2</sub> + ClONO<sub>2</sub> + BrONO<sub>2</sub>, plays an important role in the ozone depletion (NO, catalytic cycle, Crutzen 1970). At the Jungfraujoch observatory, FTIR spectrometers measure since 1984 the four most abundant members of NO<sub>y</sub>, i.e. NO, NO<sub>2</sub>, HNO<sub>3</sub> and ClONO<sub>2</sub>. Their sum is a good proxy of NO<sub>y</sub> (the most important missing gas being N<sub>2</sub>O<sub>5</sub>).

P. Demoulin<sup>(1)</sup>, E. Mahieu<sup>(1)</sup>, C. Servais<sup>(1)</sup>, B. Lejeune<sup>(1)</sup>, W. Bader<sup>(1)</sup>, G. Roland<sup>(1)</sup>, R. Zander<sup>(1)</sup>, K. Walker<sup>(2)</sup>, P. Bernath<sup>(3)</sup>, M. Van Roozendael<sup>(4)</sup>, F. Hendrick<sup>(4)</sup>.

<sup>(1)</sup>GIRPAS, Institute of Astrophysics and Geophysics, University of Liège, Belgium

<sup>(2)</sup>Department of Chemistry, University of Waterloo, Canada & Department of Physics, University of Toronto, Canada

<sup>(3)</sup>Old Dominion University, Norfolk, Virginia, USA

<sup>(4)</sup>Belgian Institute for Space Aeronomy (BIRA-IASB), Brussels, Belgium

## FTIR DATA SET

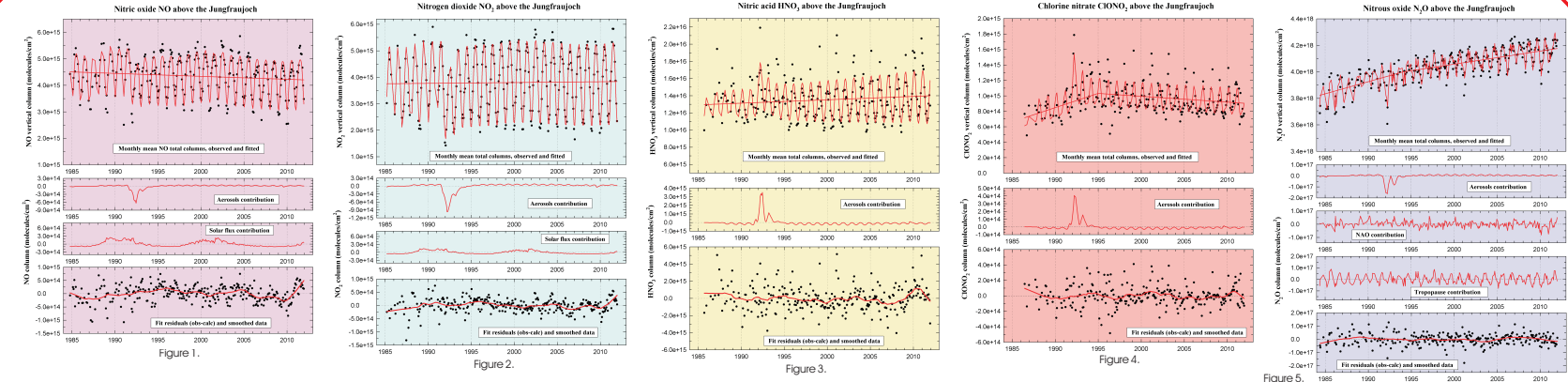
Infrared solar absorption spectra recorded since 1984 at the Jungfraujoch observatory (Swiss Alps, 3580 m a.s.l.):  
▶ 2 FTIR spectrometers (home-made and Bruker 120 HR)  
▶ high resolution (82 to 175 cm max. OPD)  
▶ only clear-sky conditions

### ANALYSIS

- ▶ total columns
- ▶ monthly means (avoid higher weight for period with many observations)
- ▶ SFIT 1 and 2
- ▶ P,T profiles from NCEP
- ▶ spectral micro-windows:
  - NO: 1899.85-1900.20, 1902.92-1903.36 and 1912.70-1912.86 cm<sup>-1</sup>
  - NO<sub>2</sub>: 2914.51-2914.86 cm<sup>-1</sup>
  - HNO<sub>3</sub>: 868.75-869.75 cm<sup>-1</sup>
  - ClONO<sub>2</sub>: 779.3-780.6 then 780.050-780.355 cm<sup>-1</sup>
  - N<sub>2</sub>O: 2481.3-2482.6, 2526.4-2528.2, 2537.85-2538.8 and 2540.1-2540.7 cm<sup>-1</sup>
- ▶ NO and NO<sub>2</sub>: empirical correction for diurnal variation

### TRENDS DERIVATION

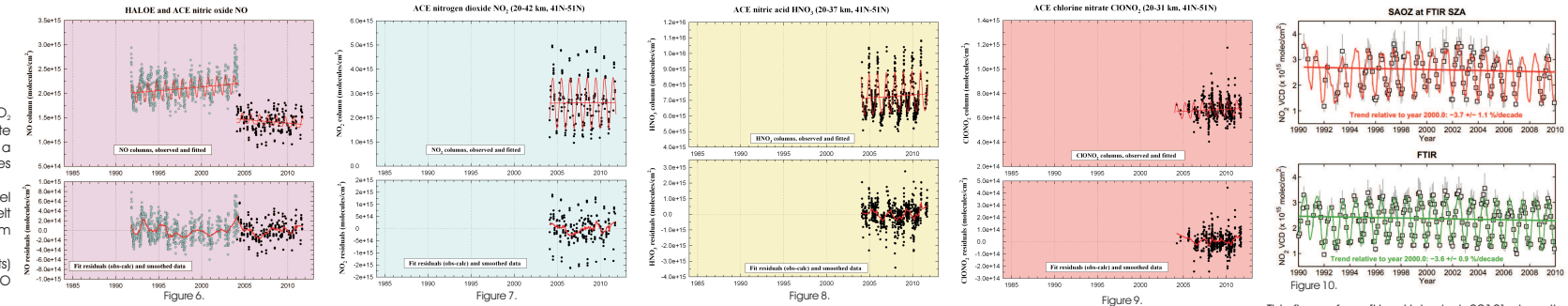
- ▶ multiple regression model, including a linear trend, a seasonal component and anomalies from various atmospheric parameters [Bodeker *et al.*, 1998]:
  - solar flux (10.7 cm wavelength, measured at Ottawa / Penticon, Canada)
  - stratospheric aerosol optical depth (15-35 km, 20° N-50° N) [Vernier *et al.*, 2011]
  - tropopause height, calculated from NCEP P,T profiles
  - other investigated parameters: NAO (North Atlantic Oscillation), QBO (Quasi-Biennial Oscillation), pressure, stratospheric temperature
- ▶ only statistically significant parameters have been kept



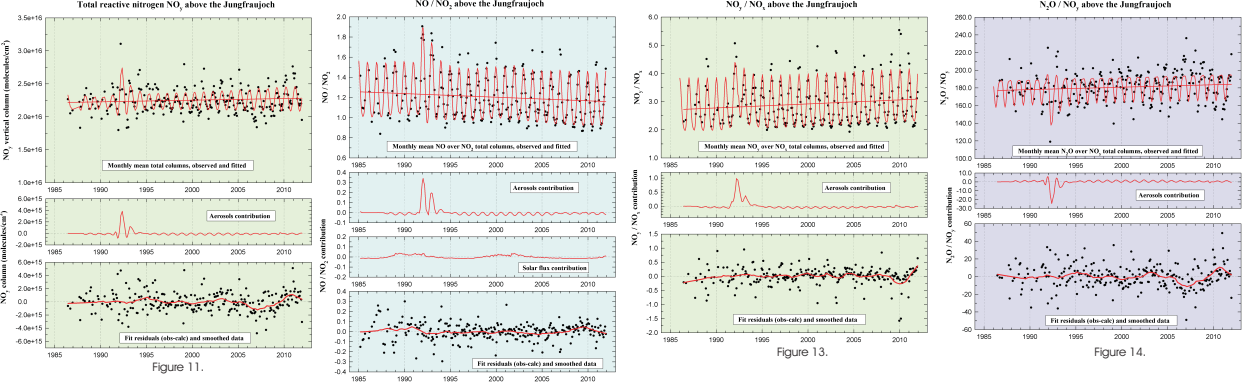
Figures 1 to 5. Upper frames: monthly mean total columns derived from Jungfraujoch FTIR spectra (black dots), together with best fit of the regression model and linear trends (red lines), respectively for NO, NO<sub>2</sub>, HNO<sub>3</sub>, ClONO<sub>2</sub> and N<sub>2</sub>O. For ClONO<sub>2</sub>, two linear trends have been included in the model, to take into account the decrease of this gas after 1995, as a consequence of the limitation of Cl emissions (Montreal protocol). Two linear trends have also been used for N<sub>2</sub>O. Middle frames give the contribution of aerosol optical depth (AOD), solar flux, NAO and tropopause (when applicable) to the total columns. The large aerosol perturbation from the Mt. Pinatubo eruption on June 15, 1991, is clearly visible, sequestering NO<sub>2</sub> into HNO<sub>3</sub>. The eleven year solar cycle influences NO and NO<sub>2</sub> total columns by a few percents. A summary of these contributions can be found in Table 1. Lower frames show the residuals of the fit (measured - model) (black dots) together with a smoothed curve (red lines) of these residuals. Note the decrease of NO, NO<sub>2</sub>, and HNO<sub>3</sub> during 2006-2007, well visible in the smoothed residuals, as well in Jungfraujoch FTS data and in ACE data. This decrease is followed by an increase in 2008, then by a new decrease in 2009 and again by a large increase in 2010.

| gas                                  | NO        | NO <sub>2</sub> | HNO <sub>3</sub> | ClONO <sub>2</sub> | NO <sub>y</sub> | NO <sub>2</sub> | N <sub>2</sub> O | NO/NO <sub>2</sub> | NO <sub>2</sub> /NO <sub>y</sub> | N <sub>2</sub> O/NO <sub>y</sub> |
|--------------------------------------|-----------|-----------------|------------------|--------------------|-----------------|-----------------|------------------|--------------------|----------------------------------|----------------------------------|
| seasonal variation (peak-to-peak, %) | 45        | 80              | 36               | 38                 | 62              | 12              | 4.7              | 47                 | 61                               | 16                               |
| month of the maximum                 | July      | July            | February         | March              | July            | October         | April            | December           | January                          | November                         |
| month of the minimum                 | January   | January         | August           | July               | January         | June            | April            | June               | August                           | April                            |
| Pinatubo max. contribution (%)       | -14       | -25             | 26               | 44                 | -37             | 17              | -4.1             | 28                 | 35                               | -14                              |
| month of maximum                     | June 1992 | April 1992      | June 1992        | March 1992         | June 1992       | May 1992        | March 1992       | Jan. 1992          | March 1992                       | May 1992                         |
| solar cycle 22 max. contribution (%) | 5.5       | 3.9             | -                | -                  | 6.9             | -               | -                | 2.9                | -                                | -                                |
| month of cycle 22 maximum            | June 1989 | Febr. 1991      | -                | -                  | June 1989       | -               | -                | Jan. 1989          | -                                | -                                |
| solar cycle 23 max. contribution (%) | 5.5       | 3.6             | -                | -                  | 5.8             | -               | -                | 3.0                | -                                | -                                |
| month of cycle 23 maximum            | Dec. 2001 | Dec. 2001       | -                | -                  | Dec. 2001       | -               | -                | Dec. 2001          | -                                | -                                |
| NAO contribution (peak-to-peak, %)   | -         | -               | -                | -                  | -               | 3.0             | -                | -                  | -                                | -                                |
| tropopause contribution (p-to-p, %)  | -         | -               | -                | -                  | -               | 3.1             | -                | -                  | -                                | -                                |

Figures 6 to 9. Upper frames: partial columns of NO, NO<sub>2</sub>, HNO<sub>3</sub>, ClONO<sub>2</sub> derived from the HALOE (blue dots) and ACE-FTS (black dots) satellite experiments. The red curves correspond to the best fits to the daily means with a linear trend and a 6-term Fourier series, to characterize the long-term changes and seasonal modulations of our target species. HALOE partial columns have been calculated above 30 mbar pressure level and for latitude belt from 42°N to 52°N. ACE partial columns are for latitude belt from 41°N to 51°N, and for altitude range of 28 - 55 km, 20 - 42 km and 20 - 31 km, respectively for NO, NO<sub>2</sub>, HNO<sub>3</sub> and ClONO<sub>2</sub>. Lower frames show the residuals of the fit (measured - model) (black dots) together with a smoothed curve (red lines) of these residuals. The minima of NO and NO<sub>2</sub> in 2007 are clearly visible in ACE-FTS data.



This figure, from [Hendrick *et al.* 2012], show the NO<sub>2</sub> vertical column time series of 2 co-located NDACC instruments: the ULg FTIR solar spectrometer and the BIRA-IASB SAOZ UV-vis instrument, both operating at the Jungfraujoch observatory. Colored lines correspond to the linear trend (thick line) and to the NO<sub>2</sub> columns recalculated using the multiple linear regression model (thin line). Trends derived from both datasets are in very good agreement.



Figures 11 to 14: same as Fig. 1 to 5, for NO<sub>y</sub> and for the ratios NO over NO<sub>y</sub>, NO<sub>2</sub> over NO<sub>y</sub>, and N<sub>2</sub>O over NO<sub>y</sub>. NO<sub>y</sub> shows no significant trend, but the minimum of 2007-2009 and the increase in 2010 are well marked. Although N<sub>2</sub>O is the source of NO<sub>y</sub>, NO<sub>y</sub> is not increasing at the same rate (0.31 ± 0.02 %/year for N<sub>2</sub>O, 0.11 ± 0.13 %/year for NO<sub>y</sub>, see Table 2). This difference is due to increasing CO<sub>2</sub> concentrations cooling the stratosphere (Rosenfield and Douglass, 1998) and to ozone and halogens changes in the stratosphere [McLinden *et al.*, 2001]. NO<sub>2</sub>/NO<sub>y</sub> is increasing at a rate of 0.53 ± 0.23 %/year, mainly due to the positive trend of HNO<sub>3</sub>, the most abundant NO<sub>y</sub> species. NO<sub>3</sub>/NO<sub>y</sub> is decreasing at a rate of -0.29 ± 0.12 %/year. This decrease is due to increased chlorine loading in the atmosphere, which increases the rate of the reaction [NO + ClO → NO<sub>2</sub> + Cl]. The increase of NO/NO<sub>y</sub> is expected in the 21<sup>st</sup> century, as a result of the decreasing chlorine loading and of CO<sub>2</sub>-induced stratospheric cooling, which slows the temperature-dependent reaction [NO + O<sub>3</sub> → NO<sub>2</sub> + O<sub>2</sub>] [Revell *et al.*, 2012].

### References

- Bodeker G. E. *et al.*, Trends and variability in vertical ozone and temperature profiles measured by ozone sondes at Lauder, New Zealand: 1986-1996. J. Geophys. Res., 103, 28661-28681, 1998.  
- Crutzen P.J., The influence of nitrogen oxides on the atmospheric ozone content. Q.J.R. Meteor. Soc., 96, 320-325, 1970.  
- Gardiner T. *et al.*, Trend analysis of greenhouse gases over Europe measured by a network of ground-based remote FTIR instruments. Atmos. Chem. Phys., 8, 6719-6727, 2008.  
- Hendrick F. *et al.*, Analysis of stratospheric NO<sub>y</sub> trends above Jungfraujoch using ground-based UV-visible, FTIR, and satellite nadir observations. Atmos. Chem. Phys. Discuss., 12, 12357-12389, 2012. doi:10.5194/acpd-12-12357-2012.  
- McLinden C. A. *et al.*, Understanding trends in stratospheric NO<sub>2</sub> and NO<sub>y</sub>. J. Geophys. Res., 106(D21), 27 787-27 793, 2001.  
- Revell L. E. *et al.*, The effectiveness of N<sub>2</sub>O in depleting stratospheric ozone. Geophys. Res. Lett., doi:10.1029/2012GL052479, in press.  
- Rosenfield J.E. and Douglass A.R., Doubled CO<sub>2</sub> effects on NO<sub>2</sub> in a coupled 2D model. Geophys. Res. Lett., 25, 4381-4384, 1998.  
- Vernier, J.-P. *et al.*, Major influence of tropical volcanic eruptions on the stratospheric aerosol layer during the last decade. Geophys. Res. Lett., 38, L12807, doi:10.1029/2011GL047563, 2011.

| gas                              | instrum. | begin  | end    | ref.   | year  | trend (%/yr) | 95 % sigma | method |
|----------------------------------|----------|--------|--------|--------|-------|--------------|------------|--------|
| NO                               | FTS      | 1984.8 | 2012.0 | 1984.0 | -0.25 | 0.12         | model & LS |        |
|                                  | FTS      | 1991.8 | 2004.2 | 1991.8 | 0.20  | 0.28         | bootstrap  |        |
|                                  | HALOE    | 1991.8 | 2004.2 | 1991.8 | 0.77  | 0.20         | bootstrap  |        |
|                                  | FTS      | 2004.1 | 2011.7 | 2004.1 | -0.73 | 0.66         | bootstrap  |        |
|                                  | ACE      | 2004.1 | 2011.7 | 2004.1 | -0.99 | 0.83         | bootstrap  |        |
| NO <sub>2</sub>                  | FTS      | 1985.2 | 2012.0 | 1984.0 | 0.08  | 0.12         | model & LS |        |
|                                  | SAOZ     | 1990.0 | 2010.0 | 1990.0 | -0.36 | 0.09         | model & LS |        |
| HNO <sub>3</sub>                 | FTS      | 1985.8 | 2012.0 | 1984.0 | 2.31  | 0.20         | model & LS |        |
|                                  | ACE      | 2004.1 | 2011.7 | 2004.1 | 2.08  | 0.94         | bootstrap  |        |
| ClONO <sub>2</sub>               | FTS      | 1985.8 | 2012.0 | 1984.0 | 0.37  | 0.27         | model & LS |        |
|                                  | ACE      | 2004.1 | 2011.7 | 2004.1 | 0.54  | 0.27         | bootstrap  |        |
| N <sub>2</sub> O                 | FTS      | 1986.5 | 2012.0 | 1984.0 | 0.08  | 0.16         | model & LS |        |
|                                  | FTS      | 1984.4 | 1994.0 | 1984.0 | 0.46  | 0.06         | model & LS |        |
| NO/NO <sub>y</sub>               | FTS      | 1985.2 | 2012.0 | 1984.0 | -0.29 | 0.12         | model & LS |        |
|                                  | FTS      | 1986.5 | 2012.0 | 1984.0 | 0.53  | 0.23         | model & LS |        |
| NO <sub>2</sub> /NO <sub>y</sub> | FTS      | 1985.2 | 2012.0 | 1984.0 | 0.16  | 0.15         | model & LS |        |
|                                  | FTS      | 1986.5 | 2012.0 | 1984.0 | 0.53  | 0.23         | model & LS |        |

Table 2. Linear trends retrieved for different gas of the NO<sub>y</sub> family and with different instruments. FTS stands for the 2 infrared Fourier transform spectrometers at the Jungfraujoch; SAOZ is the Jungfraujoch UV-visible instrument; HALOE is the space-borne Halogen Occultation Experiment aboard the UARS satellite; ACE is the ACE-FITS instrument aboard the Canadian satellite SCISAT-1. Trends are given in % per year, relatively to the years indicated in column 5. Column 7 gives the 95 % confidence level of the trends. In column 8, the method used to derive the trends is indicated, either the least squares regression model (for the longer time series) or the bootstrap resampling method [Gardiner *et al.*, 2008]. FTS and SAOZ trends are for total columns. The sensitivity of these techniques to NO<sub>y</sub> is about the same, with no sensitivity in the troposphere and a maximum sensitivity between 20 and 35 km altitude, where the NO<sub>y</sub> concentration in the stratosphere is the largest. NO<sub>y</sub> trends retrieved from these 2 techniques are in perfect agreement. HALOE trend for NO is for partial columns above 30 mbar pressure level and for latitude belt from 42°N to 52°N. NO trends retrieved from FTS data and from HALOE do not agree, probably because of different altitude sensitivity and range. ACE trends are for latitude belt from 41°N to 51°N, and for altitude range of 28 - 55 km, 20 - 42 km, 20 - 37 km and 20 - 31 km, respectively for NO, NO<sub>2</sub>, HNO<sub>3</sub> and ClONO<sub>2</sub>. Trends retrieved from FTS and from ACE are in relatively good agreement, except for HNO<sub>3</sub>.

### Acknowledgements

- Belgian colleagues from ULg, ORB/KSB and BIRA-IASB for participation to observation campaigns  
- GAW-CH program (MeteoSwiss, Zurich); HFS Jungfraujoch  
- A3C and AGACC-II (PRODEX and SSD programs from BELSPO, Brussels)  
- the EC NORS project (FP7)  
- ACE-FITS team, Canadian Space Agency (CSA) and NSERC, Canada.